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NAVAL POSTGRADUATE SCHOOL
Monterey, California



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SURFACE DUE TO INCIDENT LASER RADIATION.

by

⑩ Michael John Milchanowski

⑪ March 1977

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Thesis Advisor: D. J. Collins

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POWER-DENSITY DISTRIBUTION BELOW THE OCEAN SURFACE DUE TO
INCIDENT LASER RADIATION

by

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Submitted in partial fulfillment of the
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ABSTRACT

The time-averaged power-density distribution below the ocean surface due to incident laser radiation is examined by means of computer simulation of the geometrical optics involved with the air/sea interface and subsequent ocean penetration by the laser beam. The effects over the entire spectra of incidence angles, wind velocities, wind directions, beam spot sizes and depths of penetration are analyzed.

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I. INTRODUCTION

The time-averaged power-density distribution below the ocean surface due to incident laser radiation is examined over the entire spectra of incidence angles, wind velocities and wind directions by means of geometrical optics.

Current interest in using laser systems for communications through the air-sea interface and for detecting submerged objects necessitates the development of prediction methods for determining the subsurface power-distribution of radiation from a laser source above the ocean. A mathematical model for the optical-communications application done by Karp in Ref. 1 is developed in terms of a radiance function related to the mutual coherence function with accurate results for incidence angles out to 45 degrees. A more general model developed by Swennen in Refs. 2-4 examines the power-density distribution below the ocean surface through a rigorous assessment of the ocean surface geometry allowing for theoretically accurate results over the entire incidence angle spectrum.

The computer code developed for this analysis of the power-density distribution below the ocean surface is an expansion of Swennen's theory[Refs. 2-4].

II. DESCRIPTION OF PHYSICAL PARAMETERS

A. OCEAN SURFACE

A geometrical representation of the ocean surface in terms of local slopes developed by Cox and Munk[Ref. 5] is the basis for the analysis of the power probability distribution below the ocean, as it is in Refs. 2-4. A review of this representation follows.

The center of symmetry of the incident laser beam on the ocean surface facet is the center of a right-handed cartesian coordinate system, point "O" in Fig. 1. The z-axis points vertically upward. The y-axis is in the horizontal plane, colinear with the projection of the center of the incident laser beam onto that plane and pointing in the direction of the laser source.

The slopes of the ocean surface facet are defined by angular parameters Alpha and Beta. The angle Beta is the angle between the line of steepest ascent of the facet and the x-y plane; Alpha is the angle between the projection of the line of steepest ascent onto the x-y plane and the y-axis.

A second cartesian coordinate system at the surface is used, as explained in Ref. 2, to simplify the computation of the slope probability function. This coordinate system is designated by primes with the y'-axis pointing towards the wind source and rotated an angle Chi from the y-axis. The

z'-axis coincides with the z-axis. The slope parameters are also primed with their transformations given by:

$$\text{Beta}' = \text{Beta}$$

$$\text{Alpha}' = \text{Alpha} - \text{Chi}$$

Below the ocean surface a depth Z at the point of observation, O' , is centered a translational transformation of the surface coordinate system. Angles μ and ν define the laser beam refracted to O' by any surface facet. The first angle, μ , is the angle between the refracted ray and the z-axis; the second, ν , is the angle between the y-axis and the vertical plane of the refracted ray.

B. LASER BEAM

A cylindrical laser beam of radius a is incident on the ocean surface wave facet at an angle ψ measured from the z-axis. The intersection of the beam and the facet is a flat elliptical surface as illustrated in Fig. 2. The semi-minor axis of the ellipse is equal to the beam radius and the semi-major axis, b , is defined by:

$$b = a * \sec \psi$$

The analytical development, as in Refs. 2-4, begins at the ocean surface facet. The laser source, aiming, atmospheric propagation, etc. are extraneous to this analysis. Power density values at the surface are not quoted because the analysis deals with the ratio of the

power density at the depth of interest, P_d , and the power density at the surface, P_{ds} .

C. MAXIMUM ANGLE OF INCIDENCE

The angle of incidence with respect to the z-axis, Ψ_i , of the laser beam is analyzed from zero degrees to a maximum of 84 degrees. At greater angles the beam will not be able to strike all the wave facets within a wavelength of the wave due to the wave height. There is also the possibility, as illustrated in Fig. 3, that a ray incident at $\Psi_i > 84$ degrees could pass through the wave, re-enter the atmosphere and then re-enter the ocean; or, depending on the local slopes, the ray could undergo a total internal reflection if the Critical Angle were exceeded. The mathematical model of Swennen[Refs. 2-4] and this thesis do not address this situation.

This limitation on the maximum angle of incidence, $\Psi_{i_{max}}$, is derived from Fig. 4 and defined by:

$$\Psi_{i_{max}} = \tan^{-1}(L/2H)$$

The significant wave height, H , and the average period, T , of the wave are given by equations (1) and (2) as defined by Pierson[Ref. 7]. The average wavelength of the wave is given by equation (3) as defined by Hill[Ref. 8].

$$H = 2.14 \times 10^{-2} * W^2 \quad (1)$$

$$T = 0.81 * (2W/g) \quad (2)$$

$$L = gT^2/2 \quad (3)$$

where

W = wind velocity

g = gravitational constant

For wind velocities of 1 to 14 m/sec the maximum angle of incidence is 84.2 degrees (± 0.4 degrees) for all wind velocities; therefore, Ψ_{\max} is limited to 84 degrees.

III. POWER-DENSITY PROBABILITY INTEGRAL

A. GENERAL EQUATION

The power-density probability distribution (P_d) at the point of observation below the ocean surface (O') due to the entire beam is given by the integral of dP_d over all of the contributing surface facets. As a function of the angular parameters (μ, ν) the integral as defined in Refs. 2-4 is given by:

$$P_d = \iint_{\mu, \nu} F(\mu, \nu) d\mu d\nu \quad (4)$$

where

$$F(\mu, \nu) = P_{ds} * T * DTF * \cos(WI) * \sec(WR) \\ * [\cos(\mu) * \cos(\mu_0) + \sin(\mu) * \sin(\mu_0) * \cos|\nu - \nu_0|] \\ * P_{x' y'}(Z', Z') * \tan(\beta) * \sec^2(\beta) * J^{-1}$$

and

P_{ds} = power density at the surface facet
 T = Fresnel's transmittance coefficient
 DTF = diffuse transmittance function
 WI = angle of incidence

WR = angle of refraction

μ_0 = first quadrant angle the refracted center

ray of the beam makes with the positive
z-axis

ν_0 = positive angle the projection of the

refracted center ray of the beam onto the
xy-plane makes with the y-axis measured
from the latter to the former

$F(Z'_x, Z'_y)$ = time-average slope distribution
function

The development of equation (4) is described in great detail in Refs. 2-4 and will not be repeated here; however, for clarity the major components of equation (4) will be discussed.

B. FRESNEL'S TRANSMITTANCE FUNCTION

Fresnel's transmittance coefficient (T) is defined as the intensity ratio of the transmitted to the incident beam energy.

The collimated incident laser beam is composed of two polarization components, a transverse magnetic or parallel polarization ($T_{//}$) and a transverse electric or normal polarization (T_{\perp}). Resolving the refracted beam into its components ($T_{//}, T_{\perp}$) is a complex task due to rotation of the plane of incidence about all three coordinate axis over the range of integration. In order to deal with this complexity the beam is assumed to be rendered diffuse and unpolarized below the ocean surface due to scattering, propagation

direction variations and phase changes. This assumption allows for the averaging of T_1 and $T_{//}$ to obtain the value of T .

$$T = (T_1 + T_{//})/2$$

The equations for T_1 and $T_{//}$ used by Swennen[Refs. 2-4] are:

$$T_1 = [\sin(WI) * \sin(WR) / \sin(WI+WR)]^2 \quad (5)$$

$$T_{//} = [\sin(WI) * \sin(WR) / \sin(WI+WR) * \cos(WI-WR)]^2 \quad (6)$$

Equations (5) and (6) are in disagreement with those developed by Born[Ref. 9] and Powles[Ref. 10], given by:

$$T_1 = [2 * \cos(WI) * \sin(WR) / \sin(WI+WR)]^2 \quad (7)$$

$$T_{//} = [\cos(WI) * \sin(WR) // \sin(WI+WR) * \cos(WI-WR)]^2 \quad (8)$$

Equations (5) and (6) will give erroneously high values for the Fresnel transmittance function, especially at higher angles of incidence. Because of this, equations (7) and (8) are used in this modeling.

The geometry of the ray refraction through the ocean surface is illustrated in Fig. 5.

C. DIFFUSE TRANSMITTANCE FUNCTION

The diffuse transmittance function (DTF) accounts for beam attenuation below the ocean surface due to scattering and absorption.

The six-constant DTF developed by Duntley in Ref. 11 is simplified to a function of two constants; the backward scattering coefficient (BS) and the total absorption coefficient (AC). This simplification is valid because of the random scattering particle orientation in the ocean. The DTF is defined by:

$$DTF = K / [(AC+BS) * \sinh(K*Z) + K * \cosh(K*Z)]$$

where

$$K = [AC * (AC + 2BS)]^{1/2}$$

Z = depth of the point of observation

In current terminology this method of accounting for attenuation is a zero-angle forward scattering technique.

Typical values for BS and AC were used from measurements by Tyler[Ref. 12] and Duntley[Ref. 13]. These represent an average between cool and warm ocean waters in the blue-green spectrum (4800 angstrom).

$$BS = 0.065 \text{ percent}$$

$$AC = 0.044 \text{ m}^{-1}$$

D. SLOPE PROBABILITY FUNCTION

The time-average slope probability function, $P(Z'_x, Z'_y)$, is a statistical distribution of the ocean surface slopes developed by Cox and Munk[Ref. 5]. The distribution derived from the surface geometry is Gaussian, which is then altered by a Gram-Charlier series in order to account for slope skewness and peakedness caused by the wind.

$$P(Z'_x, Z'_y) = f(X_i, \text{Eta}, W)$$

where

X_i = standardized crosswind slope component

Eta = standardized upwind slope component

W = wind velocity

The development and the equation for the slope probability function are covered in detail in Refs. 2, 4, and 5. A clean ocean surface is assumed and the limits of applicability placed on the functional parameters are adhered to, namely:

$$|X_i| \leq 2.5$$

$$|\text{Eta}| \leq 2.5$$

$$W \leq 14 \text{ m/sec}$$

E. JACOBIAN

The power-density probability integral, equation (4),

was first developed over the slope components (Z'_x, Z'_y) . It was then transformed to a function of the angular parameters $(\text{Alpha}, \text{Beta})$ and then by the Jacobian (J) to a function of (Mu, Nu) .

$$\begin{aligned}
 P_d &= \iint_{Z'_x, Z'_y} dP_d \\
 &= \iint_{\alpha \beta} F(\text{Alpha}, \text{Beta}) d\text{Alpha} d\text{Beta} \\
 &= \iint_{\mu \nu} F(\text{Mu}, \text{Nu}) J^{-1} d\text{Mu} d\text{Nu}
 \end{aligned}$$

where

$$J = \partial(\text{Mu}, \text{Nu}) / \partial(\text{Alpha}, \text{Beta})$$

IV. METHODS OF SOLUTION

A. FAR ZONE

Swennen[Refs. 2-4] observed that when the depth (Z) of the point of observation is large compared to the beam cross-sectional radius at the surface ($Z/a > 100$) the integrand of equation (4) remains essentially constant during the integration. The integral can then be approximated by:

$$P_d = \iint_{\mu \nu} F(\mu, \nu) * J^{-1} d\mu d\nu$$

$$\approx F(\mu, \nu) * J^{-1} \Delta\mu \Delta\nu ; \text{ with } \mu = \mu_0 \text{ \& } \nu = \nu_0$$

The area $\Delta\mu\Delta\nu$ in the (μ, ν) plane is approximated by an ellipse of area ΔA , whose development is covered in detail in Ref. 3.

This is the Far Zone approximation and equation (4) becomes:

$$P_d = P_{ds} * T * \tan(\beta) * \sec^2 * \cos(WI)$$

$$* \sec(WR) * P(Z', P') * J^{-1} * \Delta A$$

$$\quad \quad \quad \begin{matrix} x \\ y \end{matrix}$$

The basic Far Zone computer program developed in FORTRAN language and run on the IBM 360 system solves for the ratio of the power density at the point of observation to the power density at the surface facet, P_d/P_{ds} . The main portion of the program allows for the entry of any combination of variables (wind, depth, beam radius, Psi, Chi, Nu_0 , backward scattering coefficient, absorption coefficient), and also computes the diffuse transmittance function, Fresnel's transmittance function and the slope probability function. Subroutines common to all variable entries are used to compute the angles (Alpha, Beta, WI, WR), the Jacobian and A.

It is possible to investigate the power-density probability distribution over a wide range of variables with a minimal of storage requirements and computation time due to the use of an IBM System/360 Source Library subroutine, NLNSYS, in solving for the coupled (Alpha, Beta) and (Mu, Nu) angle pairs from the simultaneous non-linear equations:

$$\cos(\mu_0) = [\cos(\Psi) + K \cos(\beta)]/n \quad (9)$$

$$\cot(\nu_0) = \cot(\alpha) - \sin(\Psi)/K \\ * \sin(\alpha) * \sin(\beta) \quad (10)$$

where

n = relative index of refraction (1.33)

$K = K(\Psi, \alpha, \beta, n)$

B. NEAR ZONE

When the depth of the point of observation is not large compared to the beam cross-sectional radius at the surface ($Z/a < 100$) the Far Zone approximation does not hold and the double integration of equation (4) must be carried out.

Swennen[Ref. 2] used Simpson's method of numerical integration to approximate equation (4). The method was subject to singularity points in the integration limits which increased the complexity and time required for the computation.

The double integration can be more accurately and relatively simply approximated by using the Gauss Quadrature system of solution, Refs. 14 and 15. The Two-Point Gauss-Legendre Quadrature method was used. This method, valid for polynomials up to degree three (equation (4) is of degree two), consists of first transforming the function $F(\mu, \nu)$ into a function $F(s, t)$ whose interval is $-1 \leq s \leq 1$ and $-1 \leq t \leq 1$ by letting:

$$\mu = [(\mu_u - \mu_l)s + \mu_u + \mu_l]/2 \quad (11)$$

$$\nu = [(\nu_u - \nu_l)t + \nu_u + \nu_l]/2 \quad (12)$$

and

$$d\mu = [(\mu_u - \mu_l)/2]ds \quad (13)$$

$$d\nu = [(\nu_u - \nu_l)/2]dt \quad (14)$$

The power-density integral

$$P_d = \int_{\mu_1}^{\mu_u} \int_{\nu_1}^{\nu_u} F(\mu, \nu) d\mu d\nu$$

is transformed into

$$P_d = [(\mu_u - \mu_1)(\nu_u - \nu_1)/4] * \int_{-1}^1 \int_{-1}^1 F(s, t) ds dt \quad (15)$$

The Two-Point Gauss-Legendre Quadrature approximation of equation (13) is given by:

$$P_d = [(\mu_u - \mu_1)(\nu_u - \nu_1)/4] * \sum_{i=0}^1 \sum_{j=0}^1 W_i * P(s_i, t_j) \quad (16)$$

where

$$W_i = 1.0 = \text{Gauss weight factors}$$

$$s_i, t_j = \pm(3)^{-1/2} = \text{Gauss roots}$$

The basic Near Zone computer program developed uses this Gauss-Legendre Quadrature method in solving the power-density integral coupled with the same subroutines as the Far Zone program. The main program computes the integration limits (μ_u, μ_1) and (ν_u, ν_1) after determining whether the surface projection of the point of observation is inside or outside the beam's horizontal cross-sectional area. The equations for the integration limits are well defined in Refs. 2-4.

The transformations described by equations (11)-(15) are then performed and the power-density is solved for using equation (16). As in the Far Zone program, the Near Zone

program can take any combination of variables as entries.

V. RESULTS

Results for oblique incidence angles ($0 < \Psi \leq 84$ degrees) were obtained for vertical plane cuts at $Nu_0 = 0$ degrees over the entire range of variables; however, valid data for vertical planar sections at $Nu_0 \neq 0$ degrees could not be obtained due to values of the standardized crosswind and upwind slope components being out of the range of applicability for the slope probability function. This problem is believed to be inherent in the method of mathematical analysis of the ocean surface used. Results for normal incidence ($\Psi = 0$ degrees) over the entire range of variables, including Nu_0 planar variations, were obtained.

The validity check carried out on the slope probability function, as discussed in Section III.D, clearly identified invalid data. Another source of invalid data occurred at higher incidence angles (Ψ) where the simultaneous solution of non-linear equations (9) and (10) resulted in an angle of incidence (WI) greater than 90 degrees while the slope probability function still indicated valid. This output was also easily identified and eliminated.

Computations with the Near Zone solution method used more computer time and storage than did the Far Zone solution method. For a computer solution using one value for each variable entry, the Near Zone used 36 per cent more computer time and 8 per cent more computer storage than did

an equivalent Far Zone solution. The higher time requirement for the Near Zone solution was due mainly to a greater compiling time requirement. The time difference between the two solution methods decreased as the number of variables that were incremented during one computer run increased.

A sample computer run with output is presented in Appendix E.

A. PRESENTATION OF RESULTS

Results are presented graphically to simplify comparisons of the wide range of variables. Vertical-planar sections of the power density probability distribution below the ocean surface cut through the center of the beam cross-section at the horizontal ocean surface at an angle μ_0 are used for the majority of the results presented.

This presentation is in rectangular coordinates with normalized power density in decibels on the ordinate and the angle μ_0 in degrees on the abscissa. The normalized power

density is equal to $10 \log_{10} [(P_d * Z^2) / P_{ds}]$. This normalization with respect to P_{ds} / Z^2 should produce a

distribution function that is independent of depth in the Far Zone regime, according to Swennen[Ref. 2-4], because the multiple slopes seen in a beam cross-section at the surface tend to render the radiation diffuse. The power density of diffuse radiation at large depths decreases as the square of the depth. When the effects of scattering and absorption

are included this independence does not occur, as will be discussed later; however, the normalization of the power density is used in all regimes for continuity and for comparisons with Refs. 2-4.

E. EFFECTS OF WIND VELOCITY

As wind velocity increases, the power density distribution spreads out and the maximum power density decreases slightly. This effect is presented in Figs. 6-11 for various angles of beam incidence with respect to the vertical ($\Psi = 0, 30, 60$ and 84 degrees). Figures 10 and 11 also show the close correlation between results obtained from the Near Zone and the Far Zone solution methods for $\Psi = 0$ and 40 degrees, respectively.

The spreading effect on the power density distribution is due to larger refraction angles (WR) caused by higher wave facet slopes that occur at higher wind velocities. The maximum power density decrease is caused by the spreading.

For an optical communication or detection system an increase in wind velocity would mean that the maximum intensity of the beam that could be focused to a point beneath the ocean surface would be reduced, but the width of receivable signal radiation or the search width would be increased.

C. EFFECTS OF INCIDENCE ANGLE

As the angle of incidence with respect to the vertical (Ψ) is increased the maximum power density decreases and

the distribution shifts away from the vertical. This effect can be seen in Figs. 6-11 and more clearly in Figs. 12 and 13. Figure 12 is a Far Zone solution at a depth of 50 meters for $\Psi = 0, 30, 60$ and 84 degrees while Fig. 13 is a Near Zone solution at a depth of 10 meters for the same values of Ψ .

The decrease in maximum power density is a result of greater reflectance at the surface facet as the angles of incidence (WI) increase with increases in Ψ . The shift in the center of the distributions away from the vertical is caused by the higher angles of refraction (WR) that occur as WI increases.

For an optical communication or detection system an increase in Ψ would be expected to lower the maximum beam intensity below the ocean surface and to shift the point of maximum intensity away from the point of water entry.

D. EFFECTS OF WIND DIRECTION

The effects of wind direction are small compared to the wind velocity effects. When the wind direction (χ) is perpendicular to the beam ($\chi = 90$ degrees) there is slightly less spreading of the power density distribution compared to $\chi = 0$ or 180 degrees. The maximum power density is essentially unaffected by variations in χ .

Figure 14 shows the effect of variations in χ for normal incidence and Fig. 15 for oblique incidence ($\Psi = 40$ degrees).

Wind direction appears to be of little concern to an optical communication or detection system as related to the

power density distribution below the ocean surface.

E. EFFECTS OF SPOT SIZE

Figures 16 and 17 show the effect of increasing the incident beam spot radius from 0.1 to 0.5 meters for $\Psi = 0$ and 40 degrees, respectively. This is done with the assumption that the energy per unit area remains constant; therefore, the total energy must increase. The power density spreads out and increases in intensity uniformly over the distribution as the spot size is increased.

Since the equations used in this solution method deal with the ratio of P_d/P_{ds} , these results simply mean that more power in a larger beam will produce a more intense and more widely distributed pattern below the ocean surface.

F. EFFECTS OF DEPTH

The decrease of power density with increases in depth is illustrated in Figs. 18 and 19 for Ψ equal to 0 and 40 degrees, respectively. The decrease is linear for the normalized power density as can be seen in Fig. 20.

Scattering and absorption cause this decrease in the power density. As the depth increases the diffuse transmittance function (DTF) rapidly becomes an inverse function of the sum of the hyperbolic sine of the depth and the hyperbolic cosine of the depth,

$$DTF = f[(\sinh Z + \cosh Z)^{-1}]$$

which means the DTF decreases rapidly with increases in depth and approaches zero in the limit. This agrees with the theory of Preisendorfer[Ref. 16].

Swennen's prediction of a constant normalized power density distribution in the Far Zone regime due to the diffusing effects of the ocean in Refs. 2-4 holds only when the effects of scattering and absorption are ignored. One computer run was made with the DTF set equal to one to illustrate this effect in Fig. 20.

Figures 21 and 22 illustrate the decrease in power density with depth, non-dimensionalized with respect to spot size, Z/a , again for Ψ equal to 0 and 40 degrees, respectively.

An optical communication or detection system will be limited greatly by the power density decrease associated with depth.

G. MAXIMUM POWER DENSITY

The maximum obtainable power density available over the spectrum of incidence angles ($\Psi = 0 - 84$ degrees) for a given set of conditions is presented in Figs. 23 and 24. The maximum power density obtainable does not occur for normal incidence, but for a small value of Ψ , because of the fact that as the wind increases the most probable slope is not zero but a small angle. The maximum obtainable power density then drops continuously with increases in Ψ after the peak that occurs around $\Psi = 10 - 15$ degrees.

Figure 23 illustrates the decrease in the maximum obtainable power density for all values of Ψ as the wind velocity increases. This is due to the spreading effect on the power density distribution caused by the wind that was described previously.

Figure 24 illustrates the independence of the maximum power density on the wind direction (χ) for all values of Ψ . χ equal to 270 degrees is not plotted but is equivalent to the condition at 90 degrees because of the symmetry of the slope probability function about the wind direction.

The angle with respect to the vertical from the point of observation below the ocean surface (μ_0) at which the maximum power density occurs for each value of Ψ is plotted in Fig. 25. The plot is independent of wind velocity and direction. Examining Figs. 16 - 19 it can be seen that Fig. 25 is also independent of depth and spot size. μ_0 for maximum power density increases nearly linearly from approximately 1 to 56 degrees as Ψ goes from 0 to 84 degrees. The increase in μ_0 is due to the power density distribution shifting away from the vertical for increasing values of Ψ as discussed earlier and illustrated in Figs. 12 and 13.

The obvious consideration with an optical communication or detection system is the decrease in maximum power density that accompanies higher incidence angles and wind velocities. The strong independence of the angle μ_0 at which the maximum power density occurs with respect to the wind velocity, wind direction, depth and spot size indicates

that a satellite based laser navigation system for submarine usage may be a possibility.

VI. CONCLUSIONS

The time-averaged power-density distribution below the ocean surface due to incident laser radiation has been examined through computer simulations in both the Near Zone and the Far Zone regimes. The effects of altering incidence angles, wind velocities, wind directions, beam spot sizes and depths of penetration have been presented.

The power-density distribution was found to be highly dependent on the angle of beam incidence, the wind velocity and the attenuation associated with depth of penetration of the laser beam. The fact that the location of the maximum power density was found to be fundamentally dependent on only the angle of beam incidence indicates a possibility of direction finding capabilities by a submerged receiver.

Improvements in computer systems and analysis techniques have allowed an examination of a much broader range of variables than were possible heretofore. Angles of beam incidence up to 84 degrees were considered within the range of validity for restrictions placed on the simulation. Above 84 degrees incidence angle the possible interactions at the air/sea interface make an analysis of the power-density distribution extremely complex, and experimentation will be necessary to give indications of the feasibility of laser operations in this area.

This simulation could be extended to an analysis of much larger spot sizes that would be expected from a satellite based transmitter. Prettyman and Cernak[Ref. 17] suggested that the effects of larger spots could be approximated by

assuming each wave facet independent of its neighbors and then integrating over the larger spot area, since the small spot represents the outer limit of the effects of the ocean surface on a laser beam of greater spot size. There would have to be a large number of surface facets contained in the spot to insure facet independence and accurate approximations, hence a strong dependence on the wavelength of the wave associated with wind velocity.

An extension of the simulation to pulsed laser radiation is also warranted, since this mode of operation may be required for air-to-subsurface laser systems. The effects of pulsed radiation occur nearly instantaneously as compared to the time-averaged analysis performed in this simulation.

A further refinement of the simulation would involve a more complex modeling of the beam attenuation due to scattering and absorption in the ocean through the use of a diffusion or a multiple forward scattering method, both described by Bucher[Ref. 18] and Preisendorfer[Ref. 16].

APPENDIX A

FIGURES

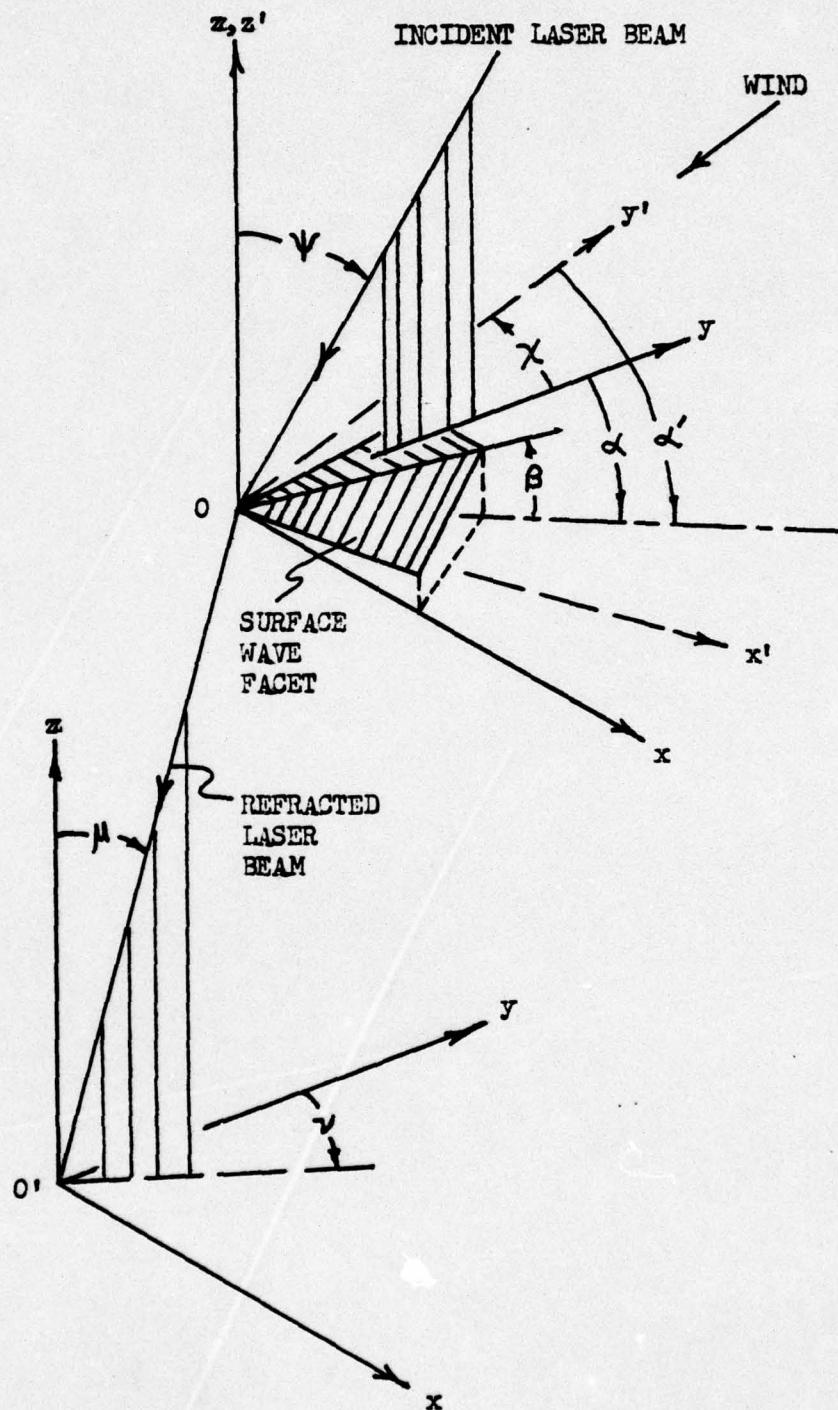


Figure 1: Coordinate systems

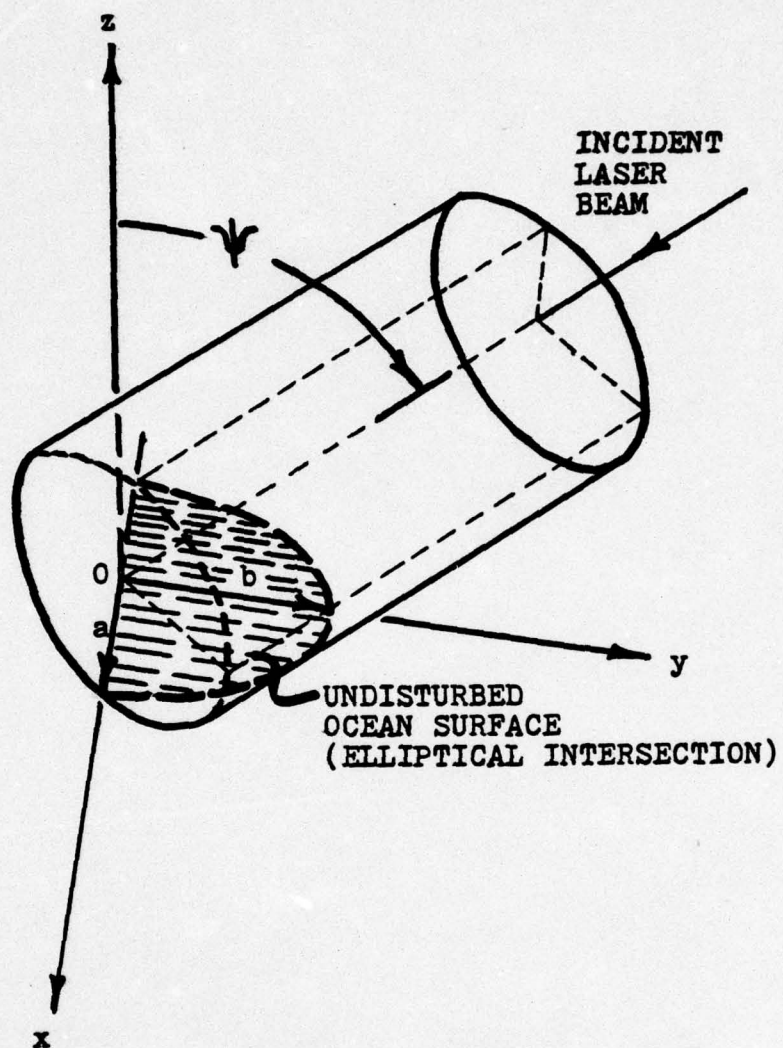


Figure 2: Intersection of laser beam
with ocean wave facet

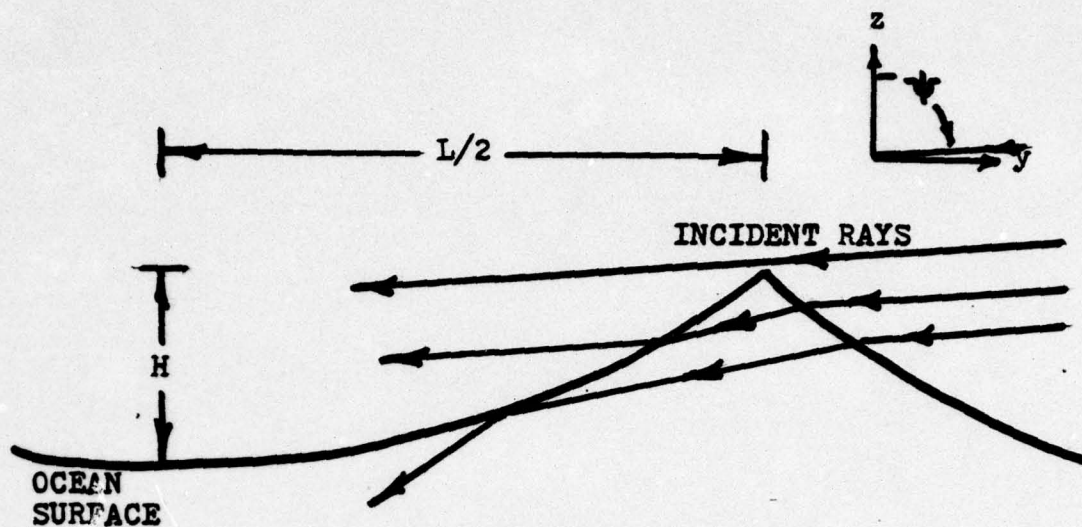


Figure 3: Possible beam paths at high incidence angles

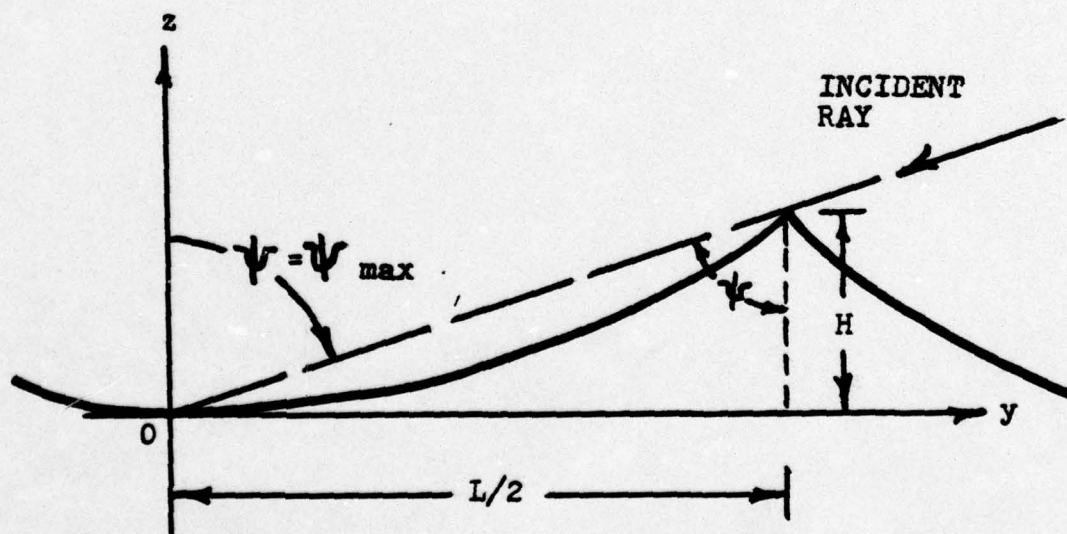


Figure 4: Maximum angle of incidence wave geometry

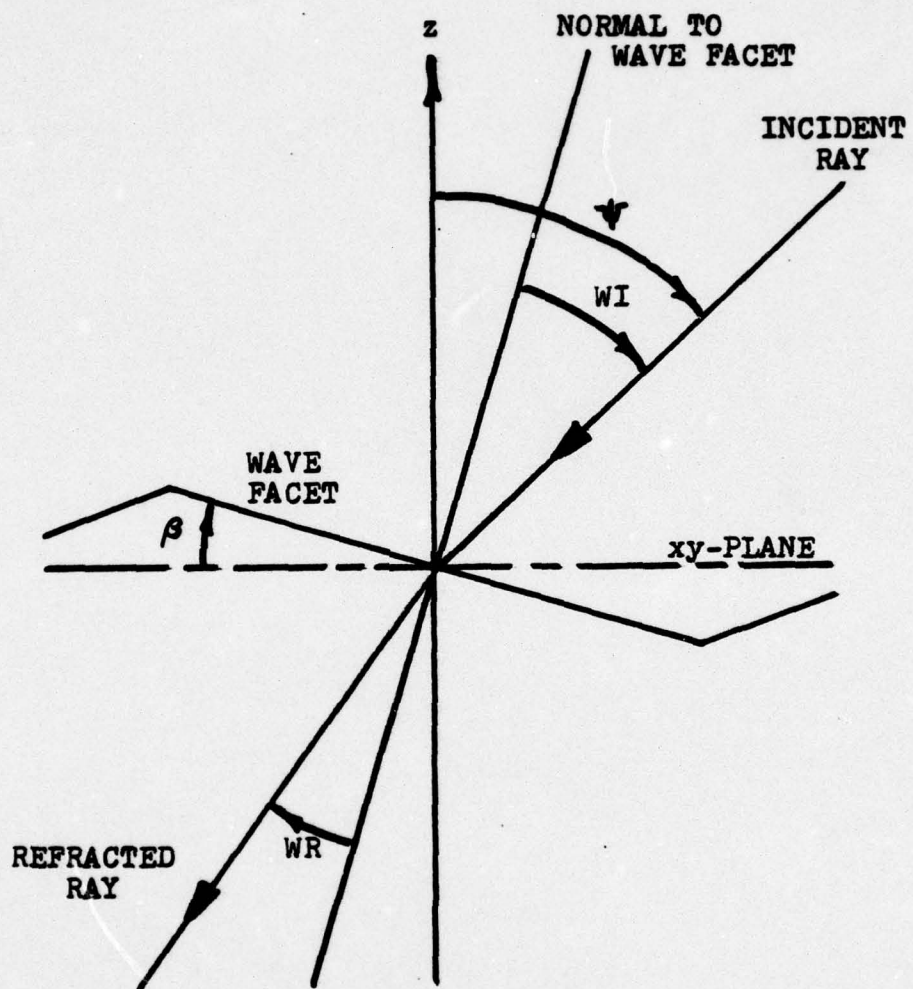


Figure 5: Ray refraction geometry

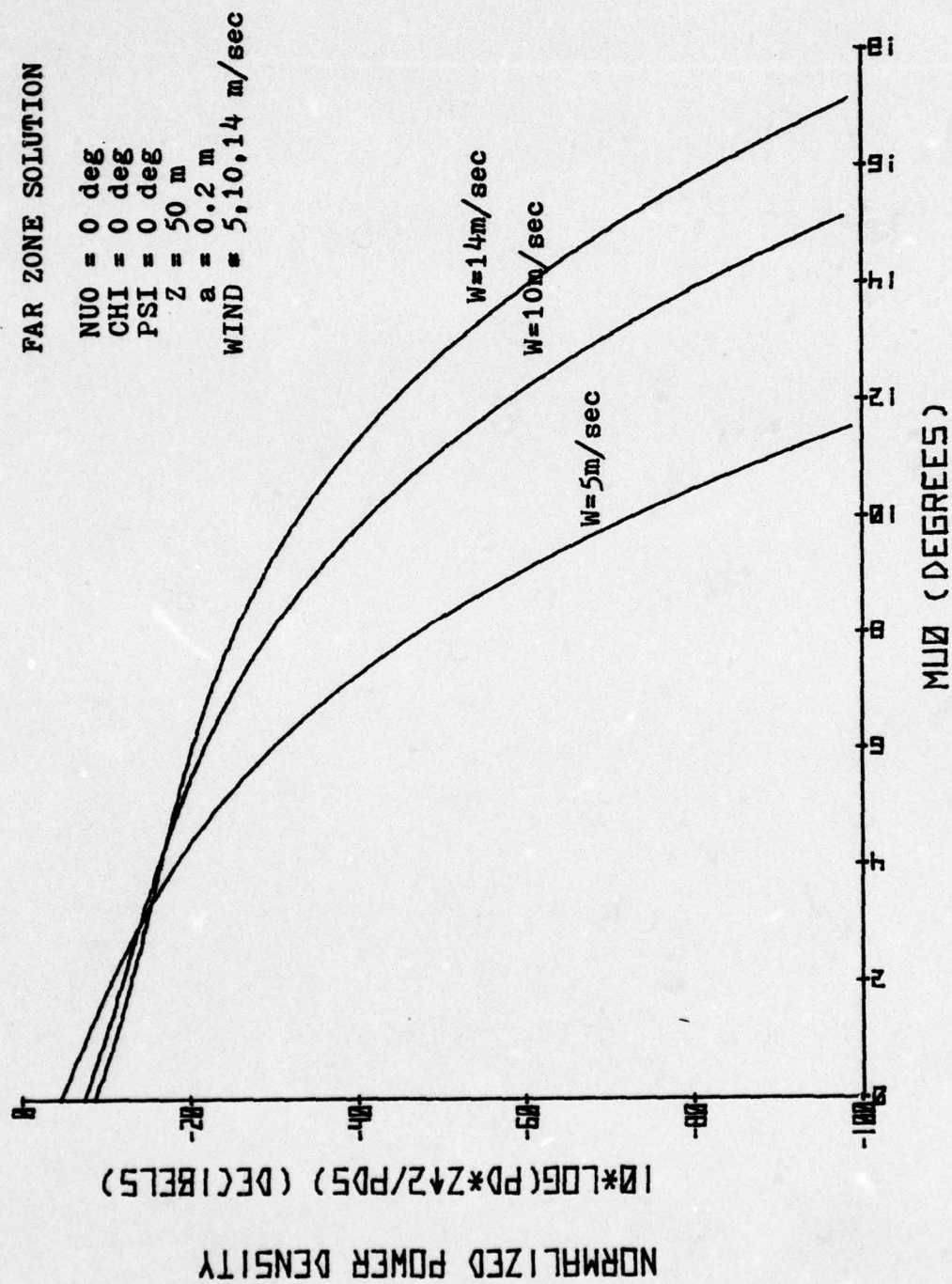


Figure 6

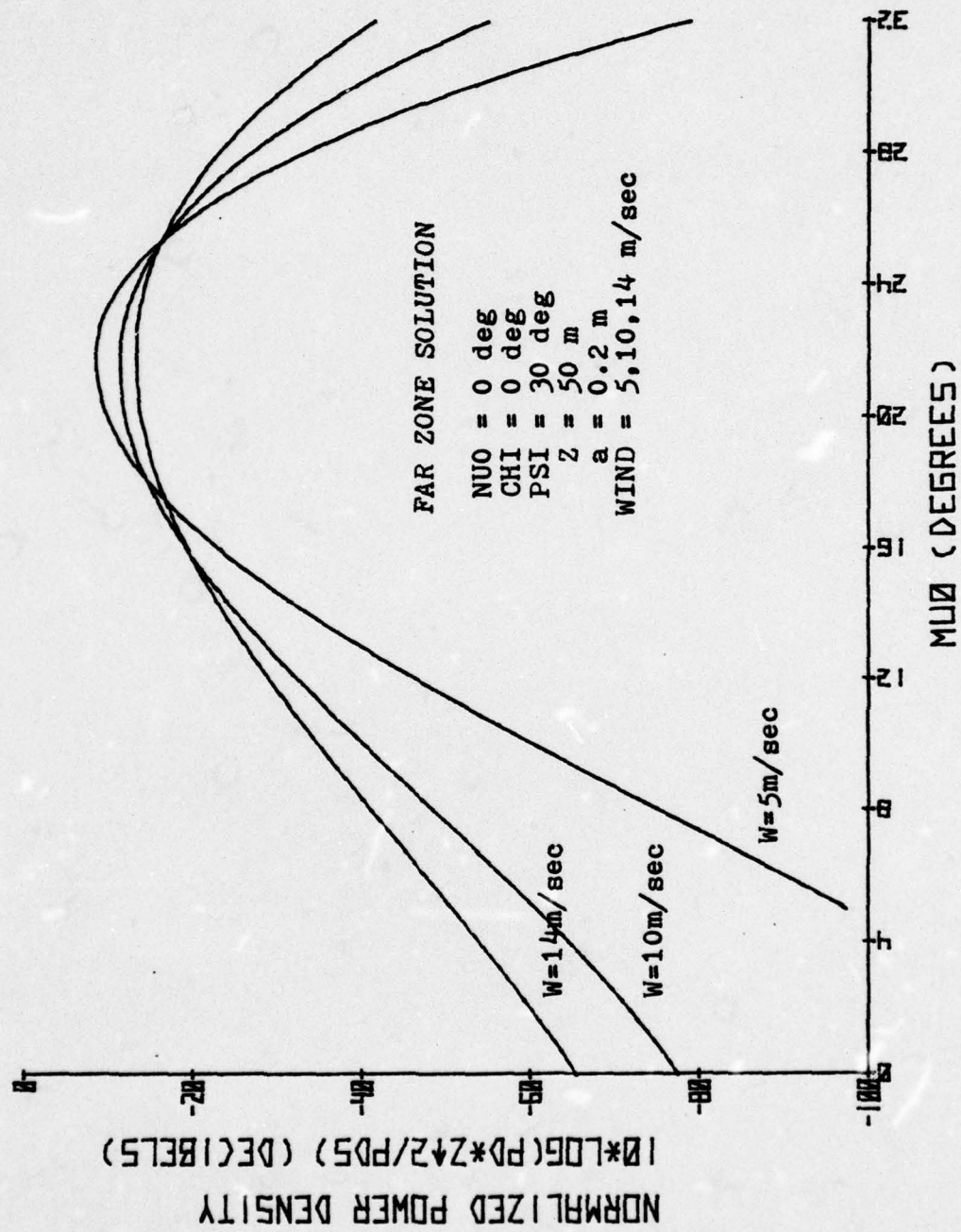


Figure 7

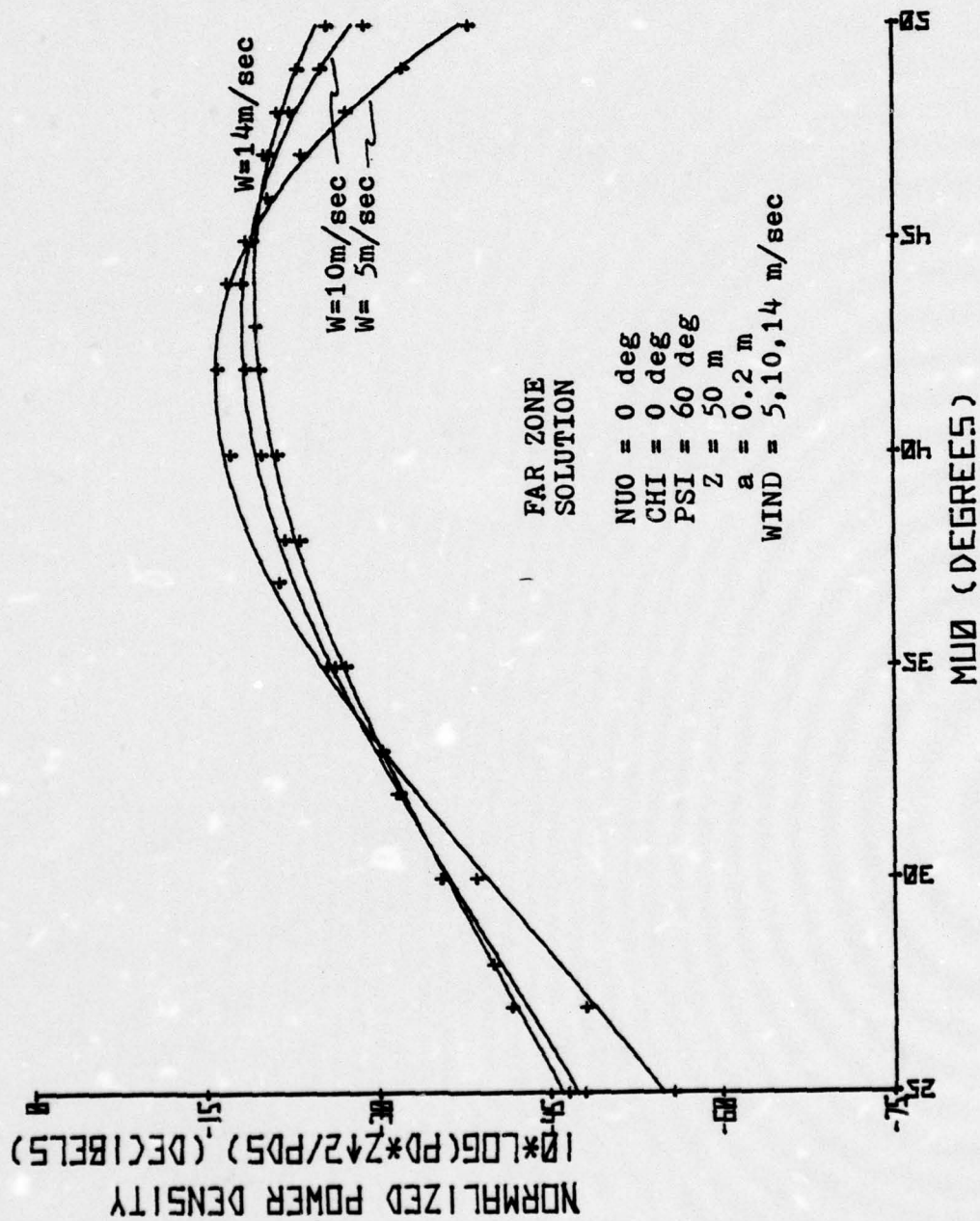


Figure 8

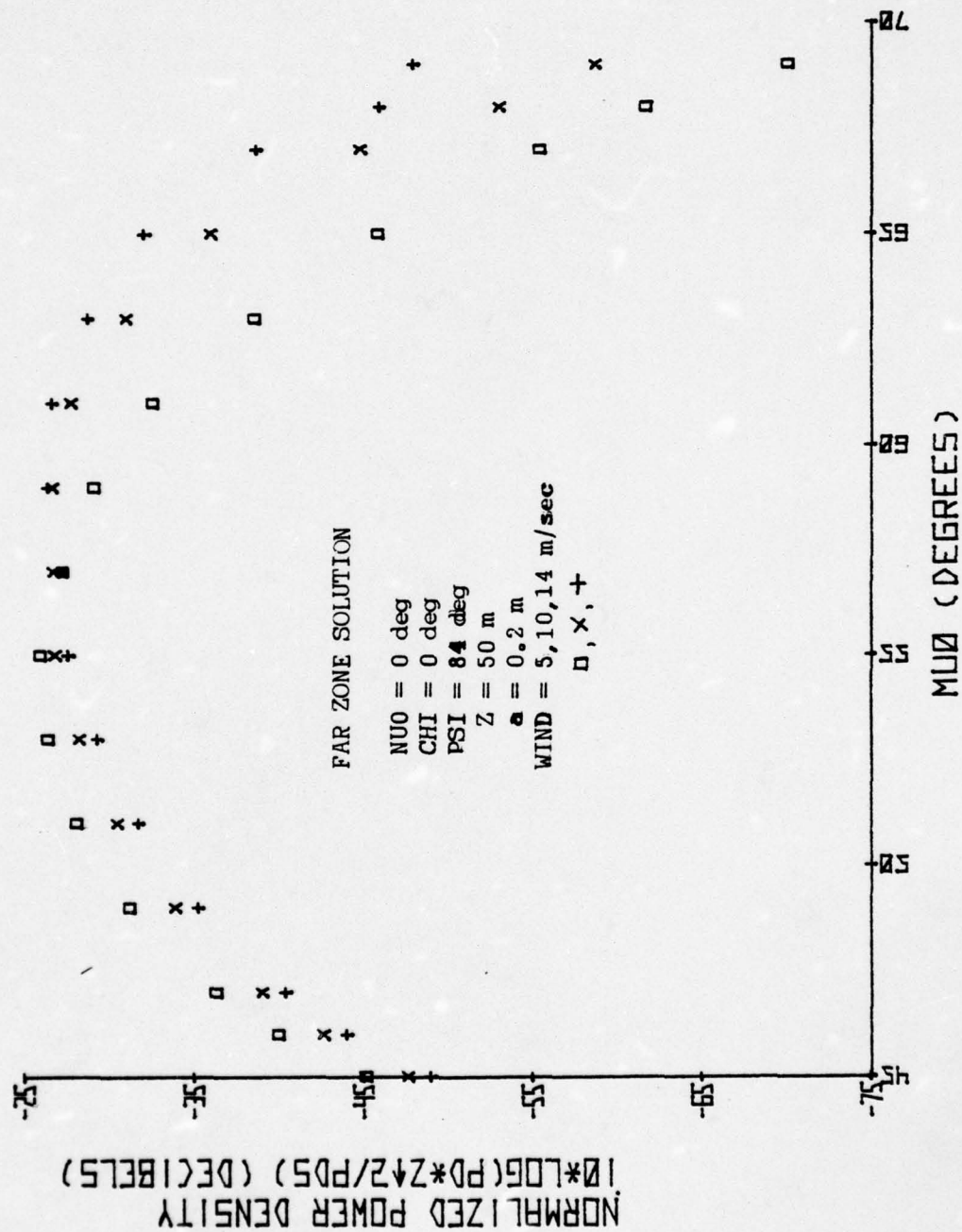


Figure 9

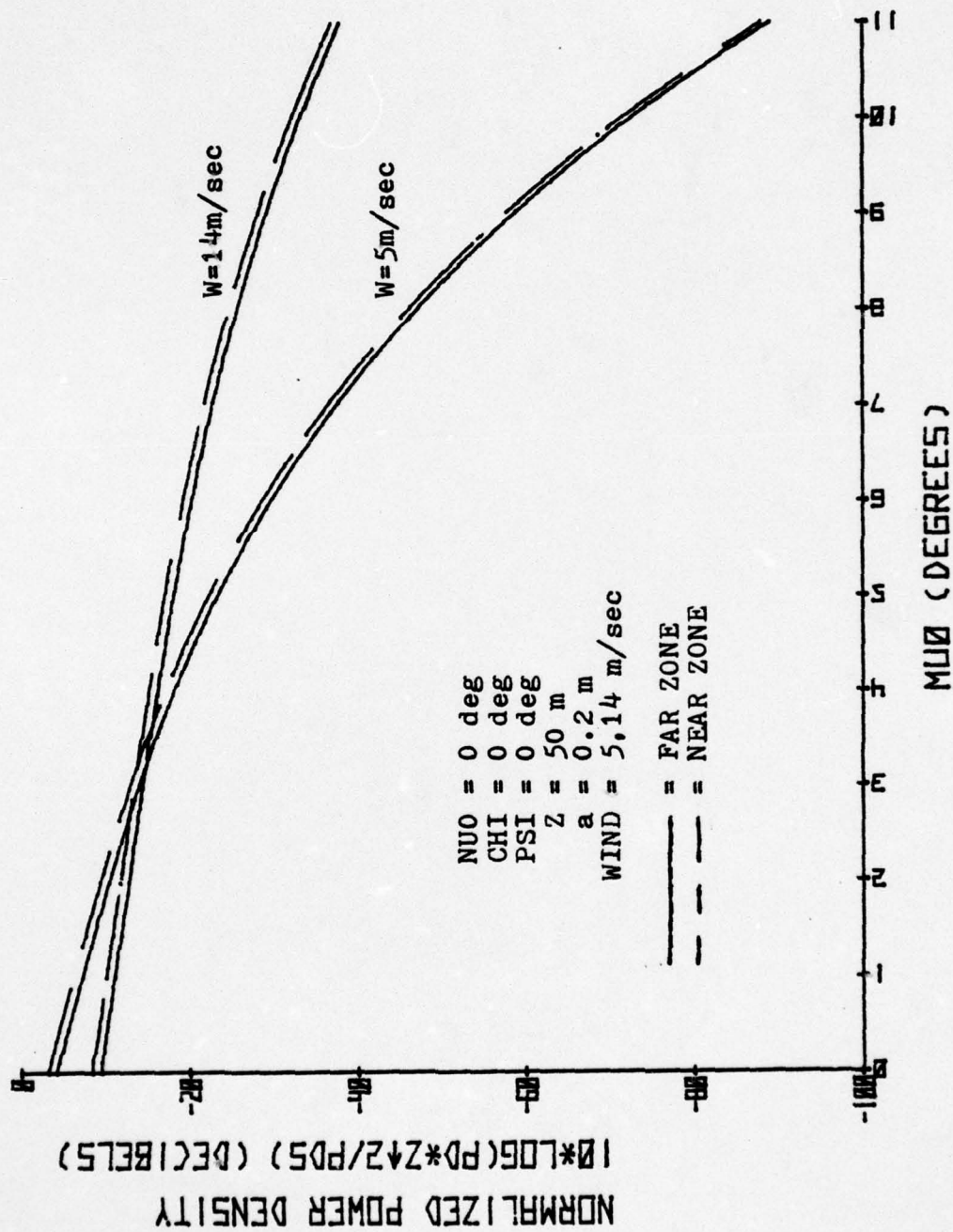


Figure 10

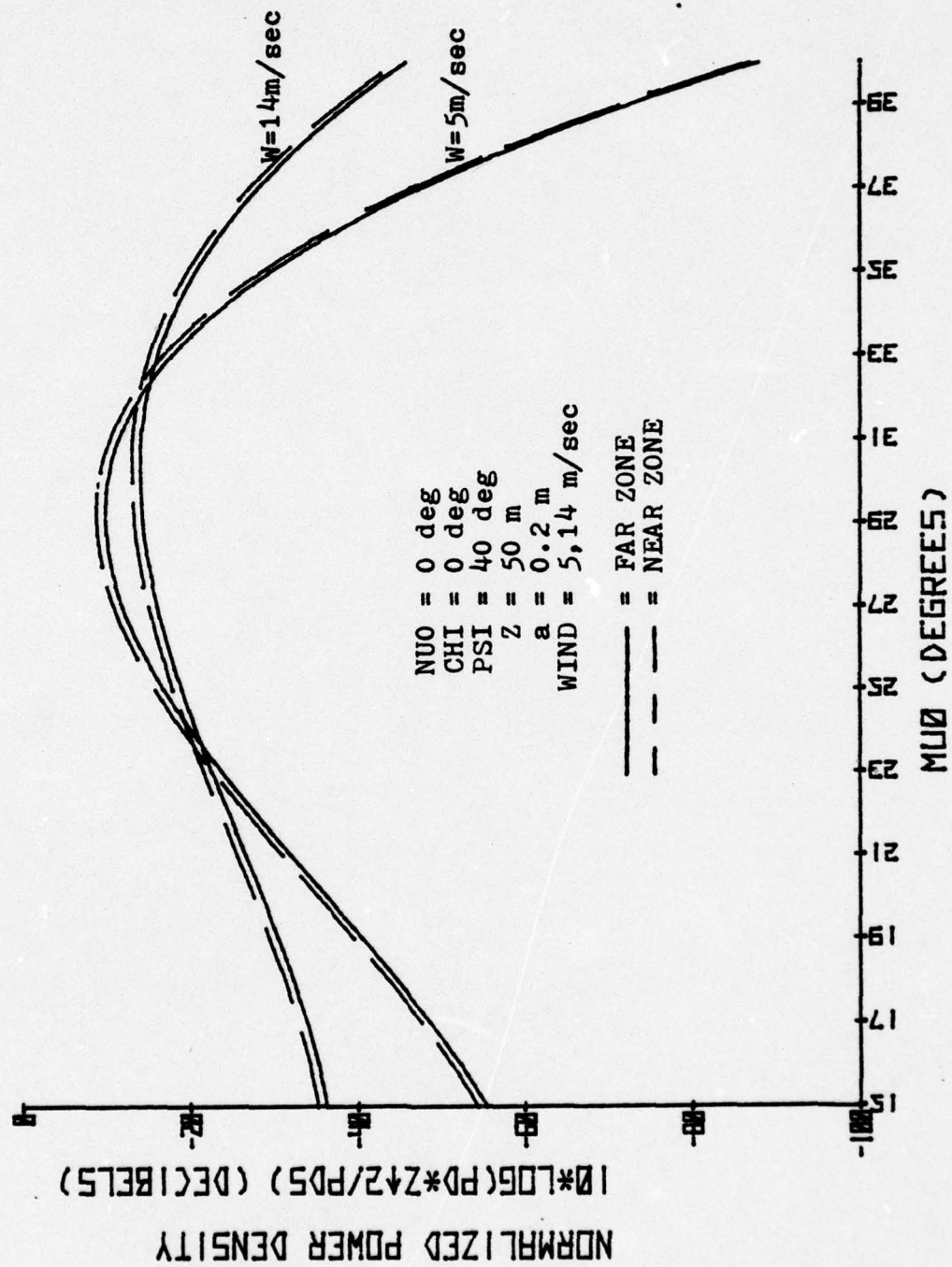


Figure 11

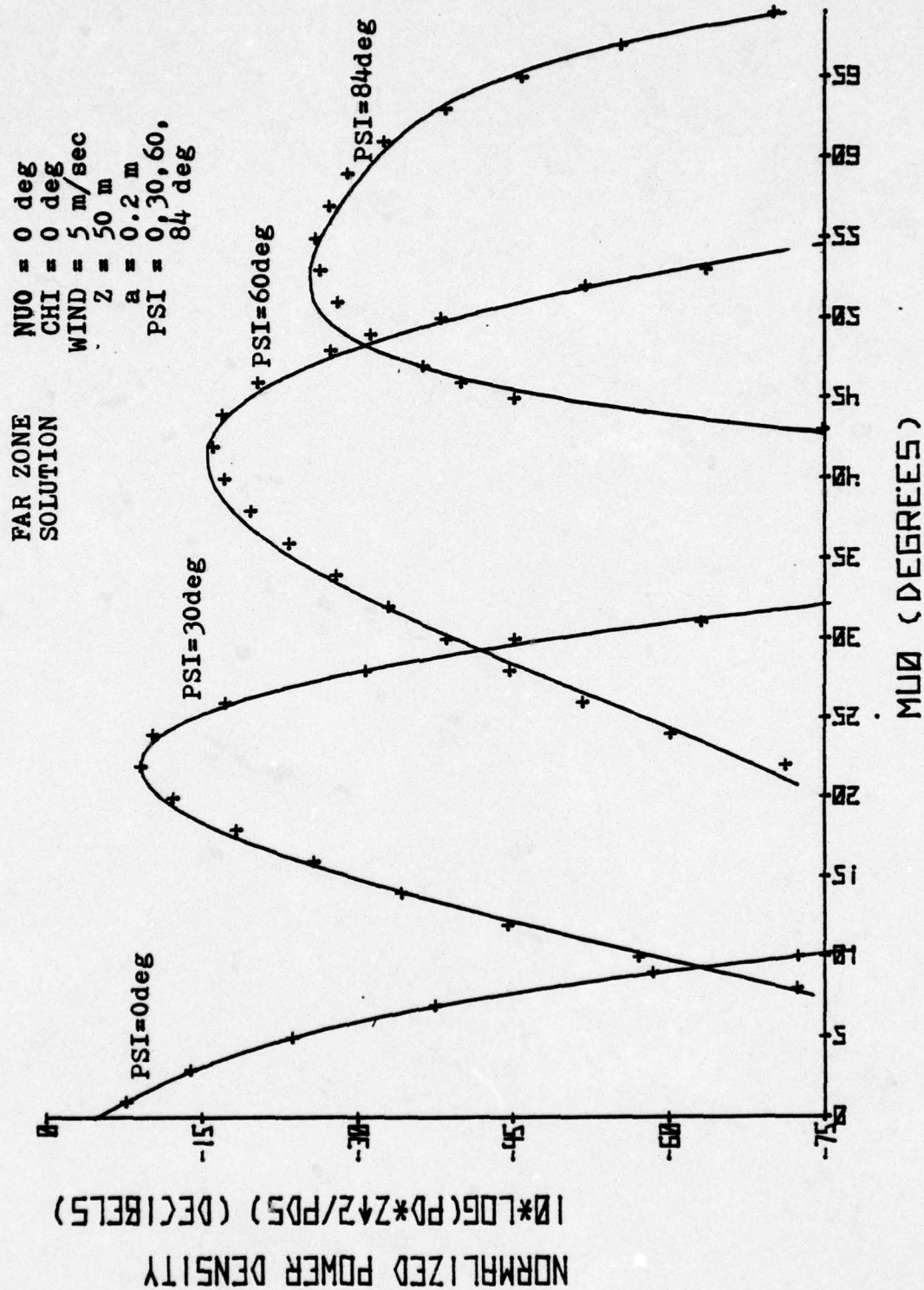


Figure 12

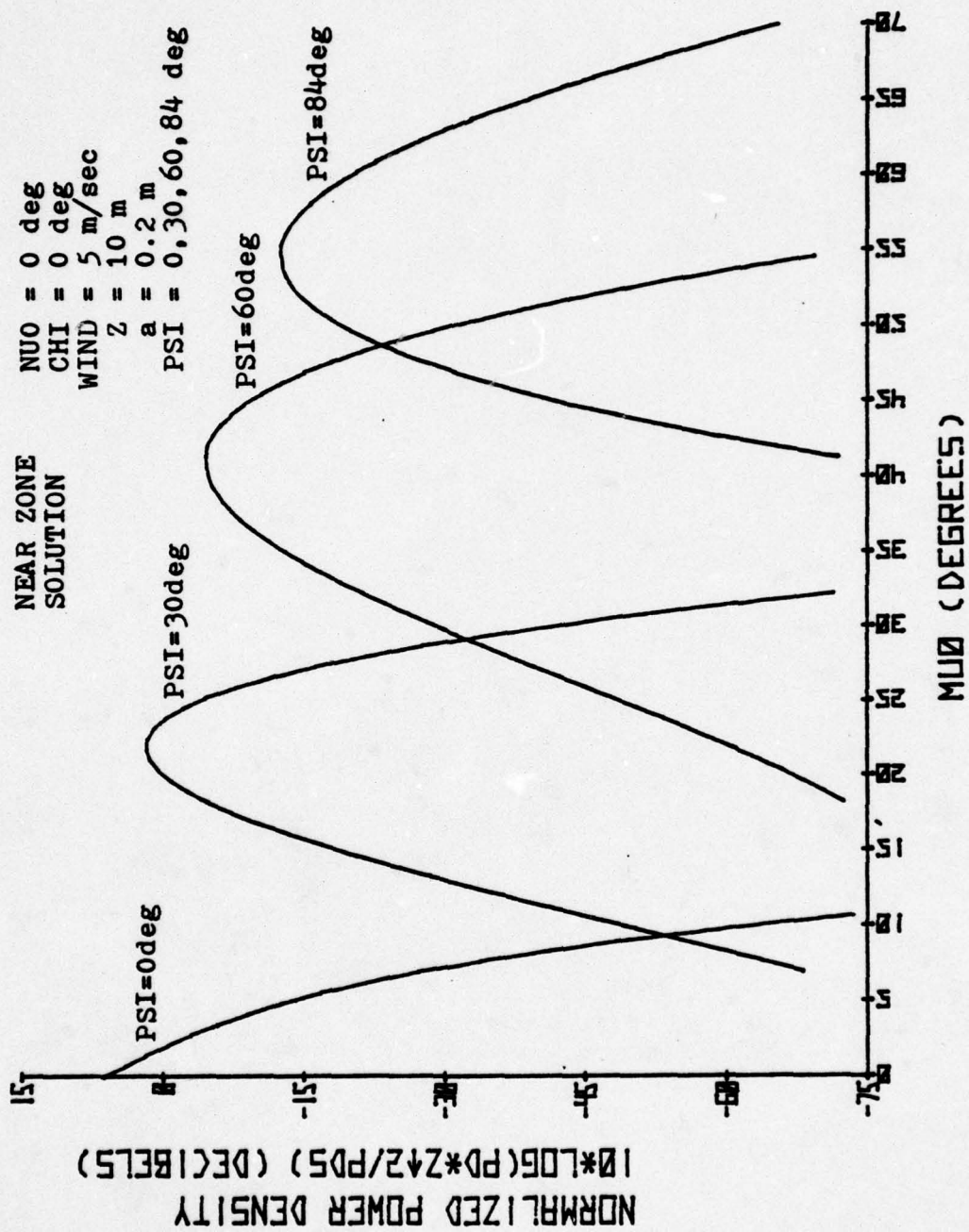


Figure 13

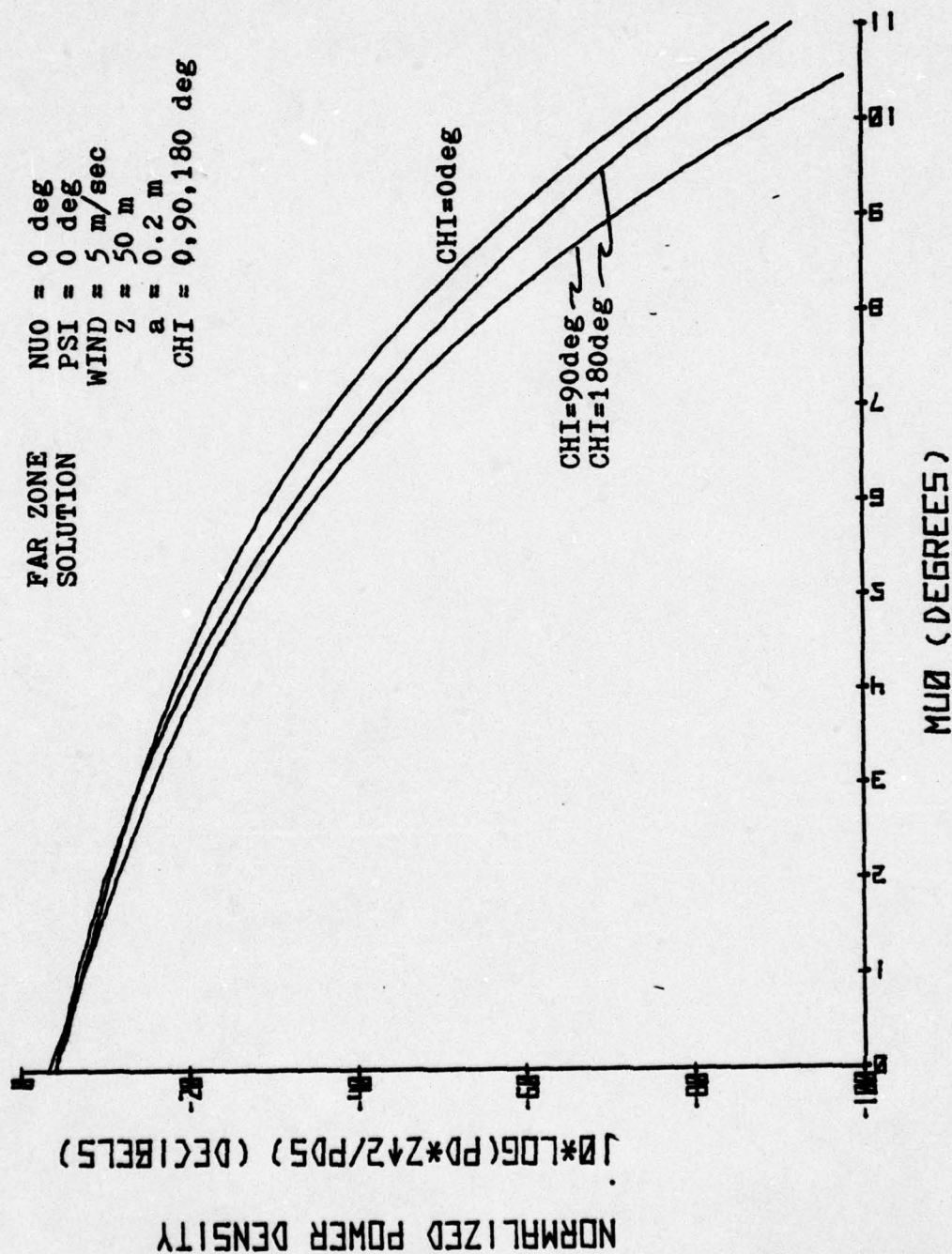


Figure 14

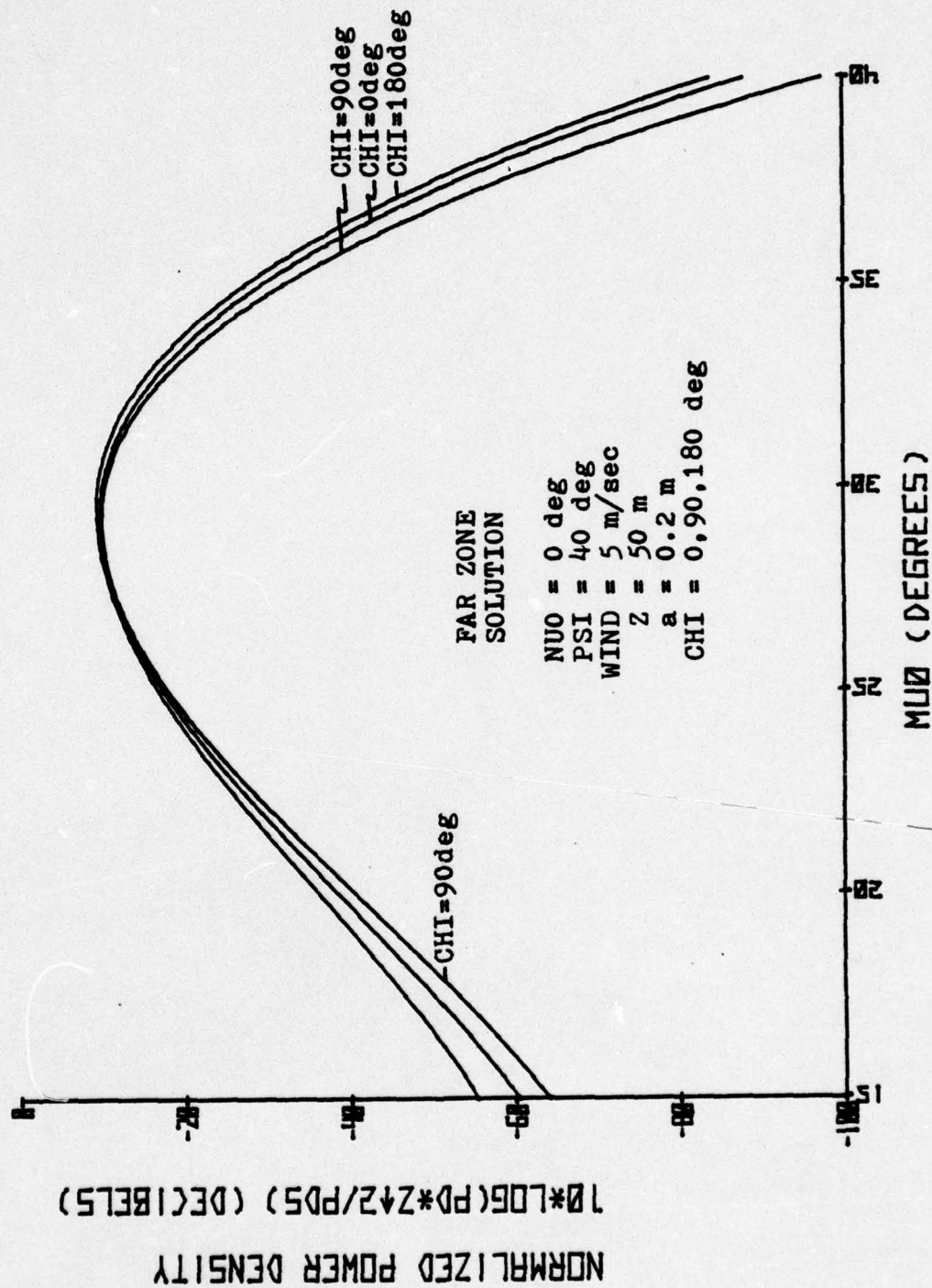


Figure 15

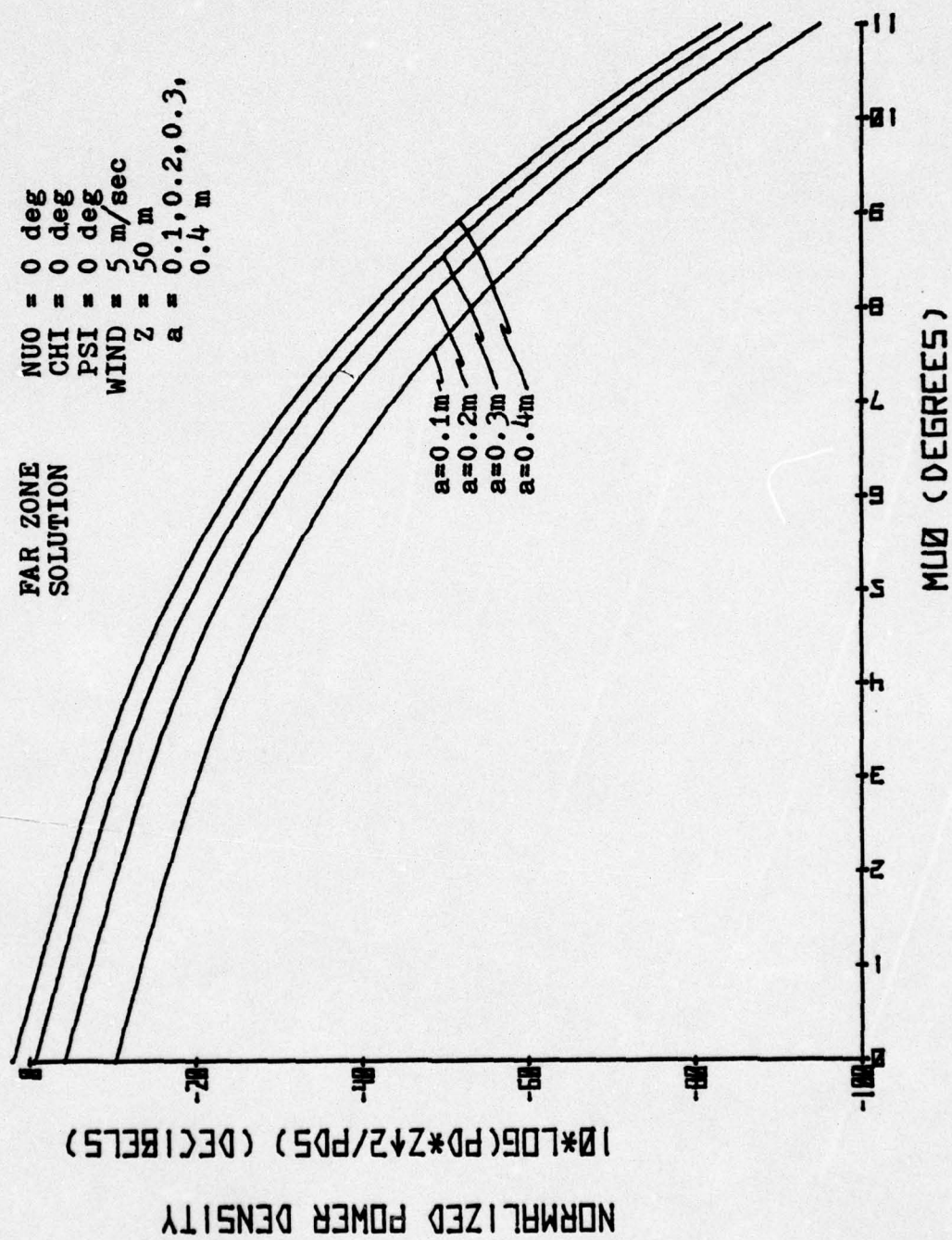


Figure 16

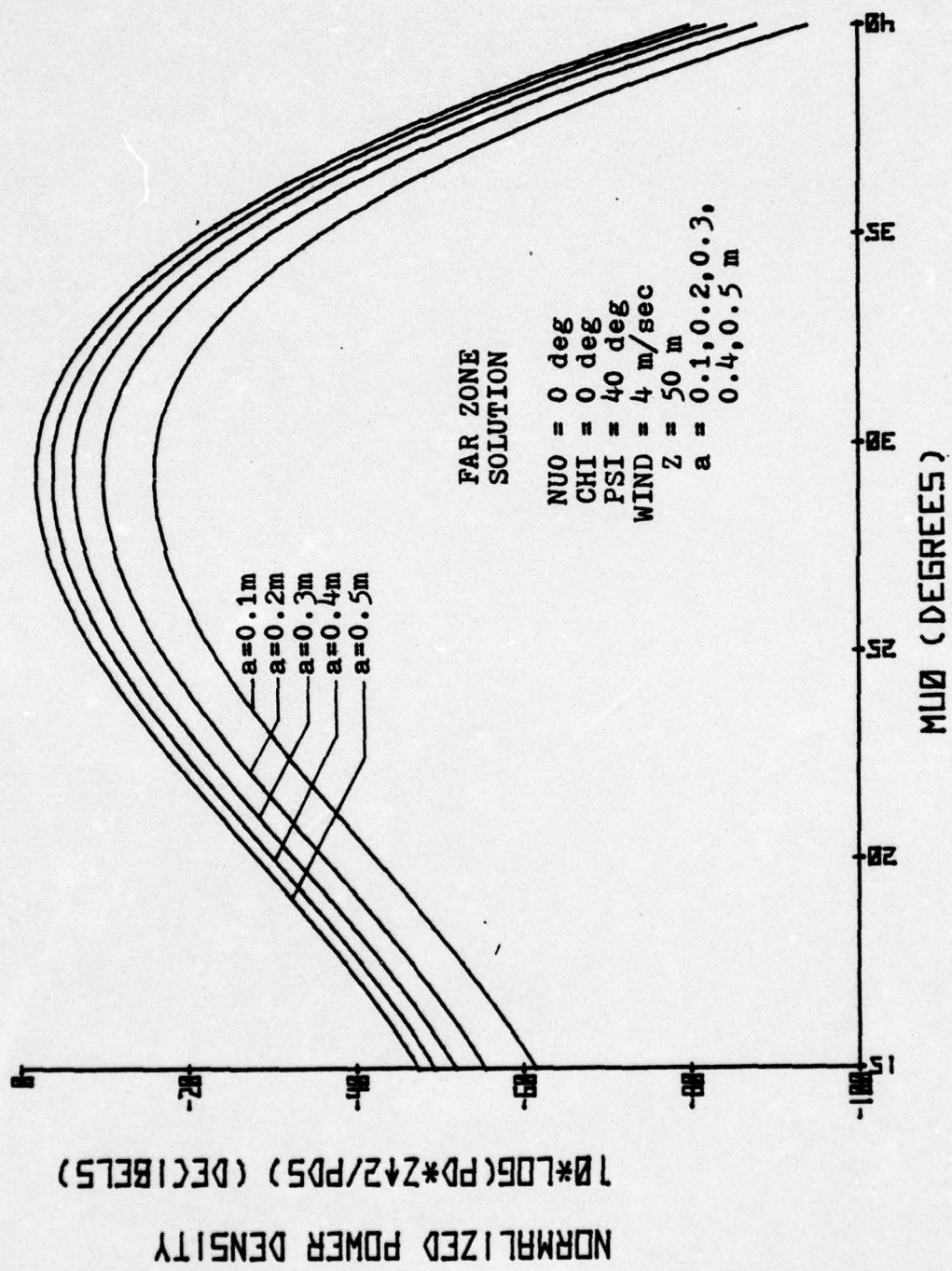


Figure 17

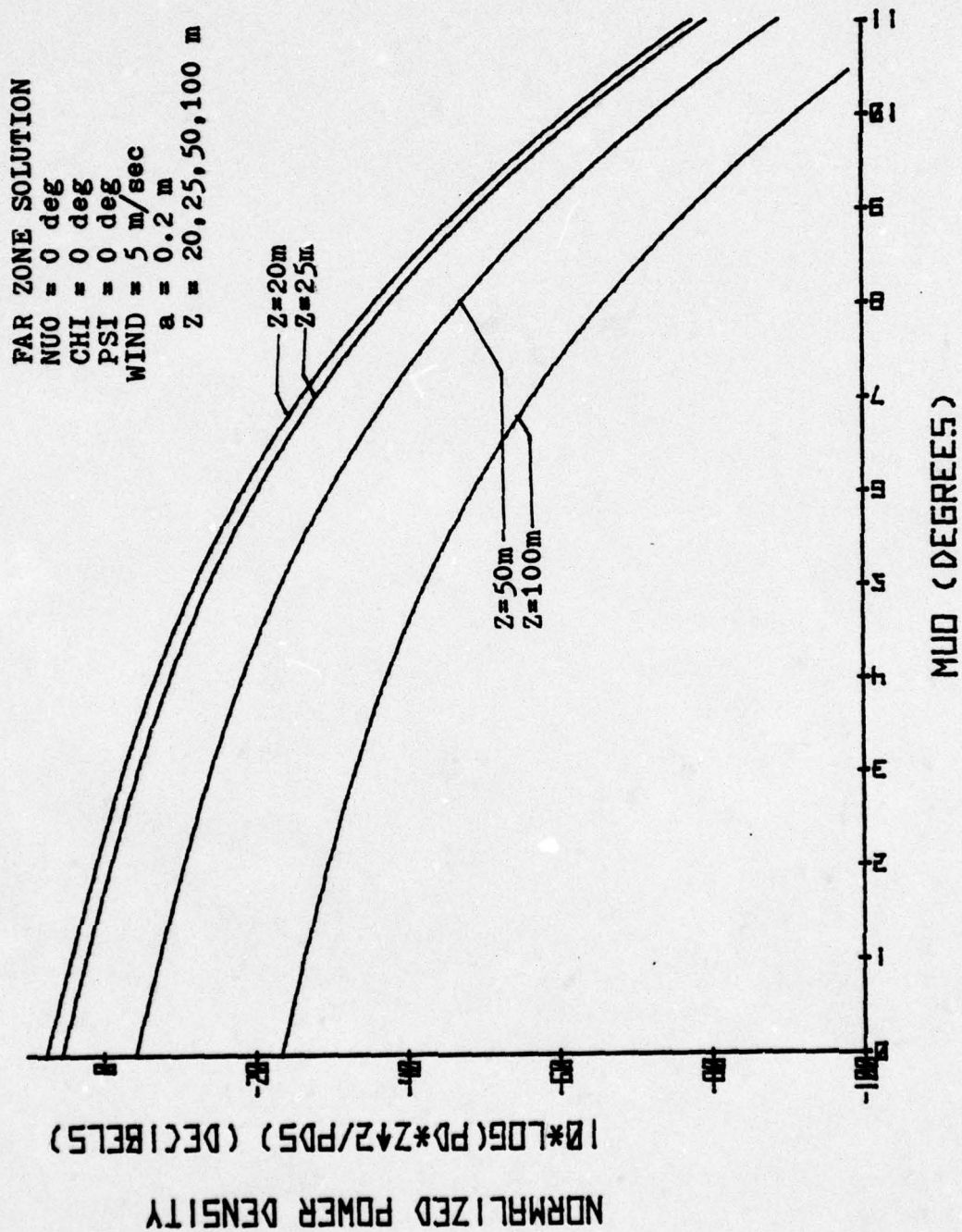


Figure 18

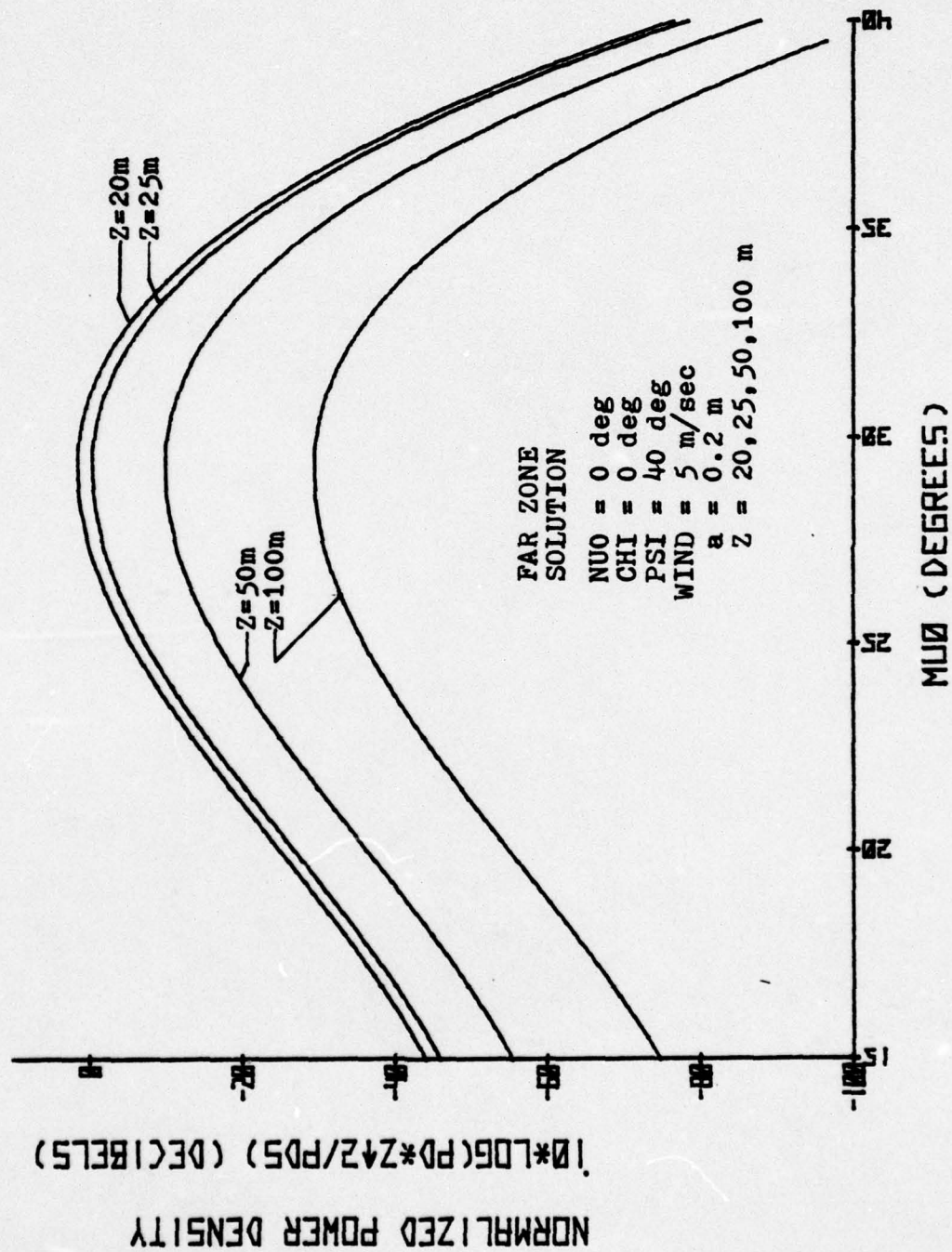


Figure 19

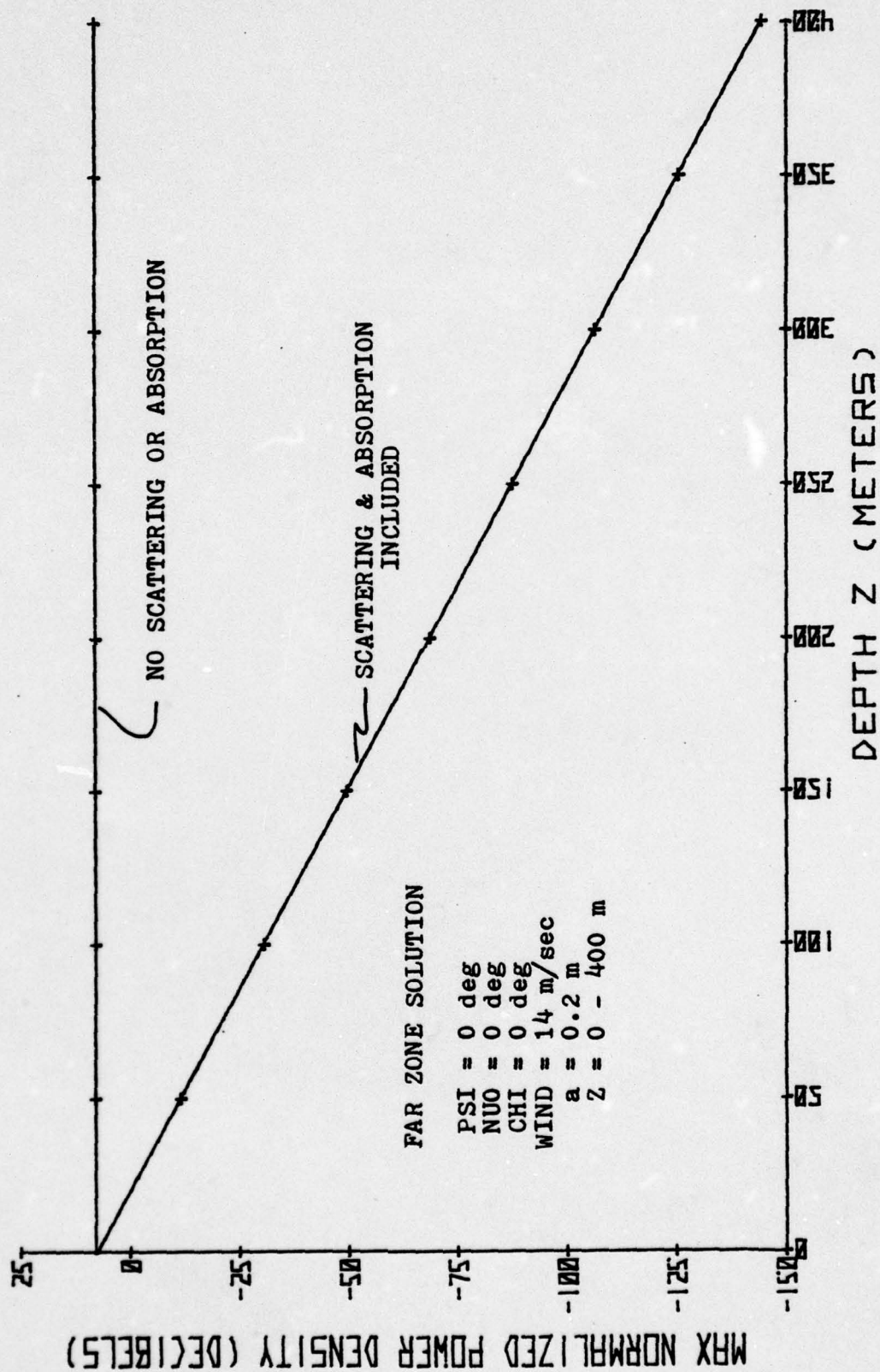


Figure 20

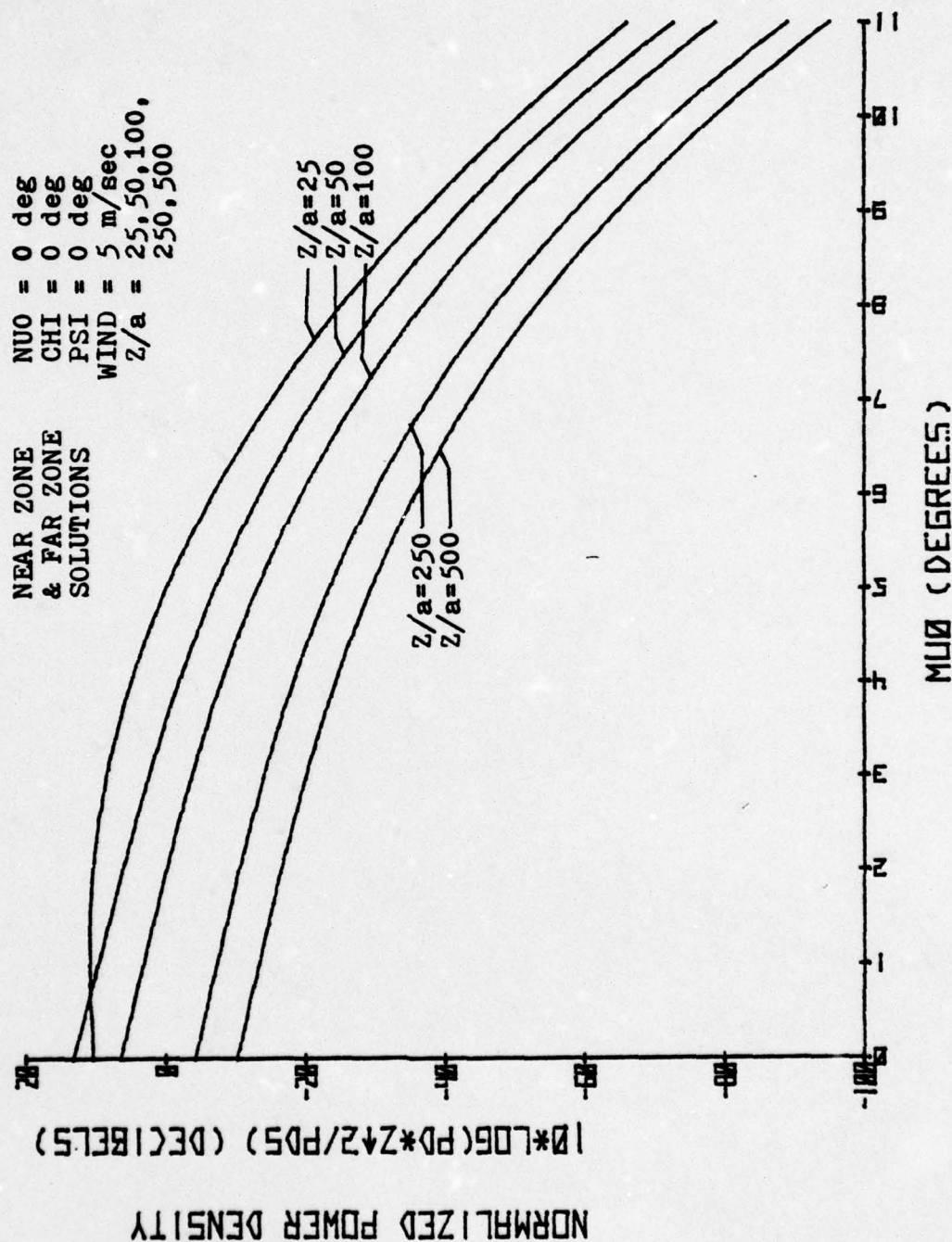


Figure 21

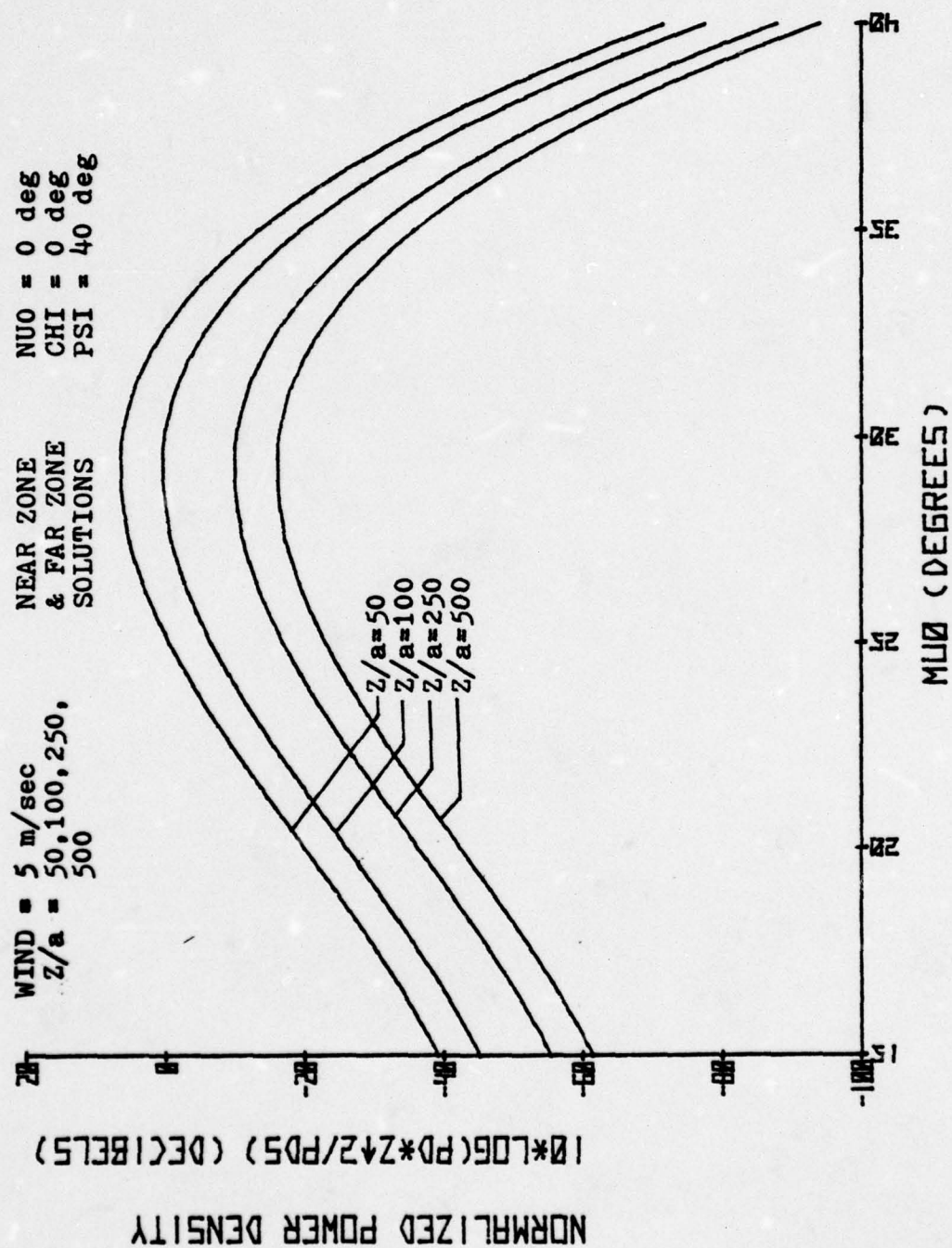


Figure 22

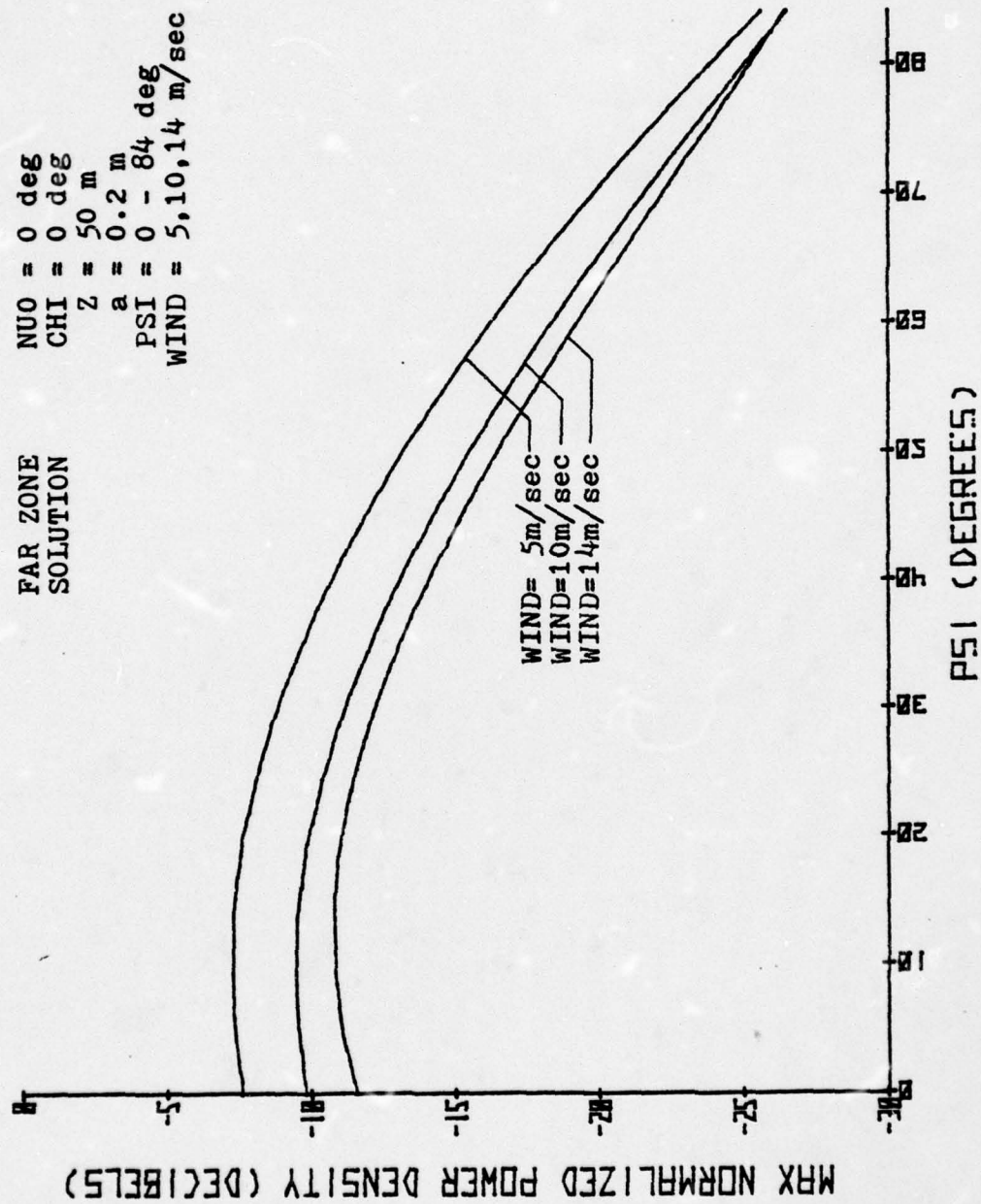


Figure 23

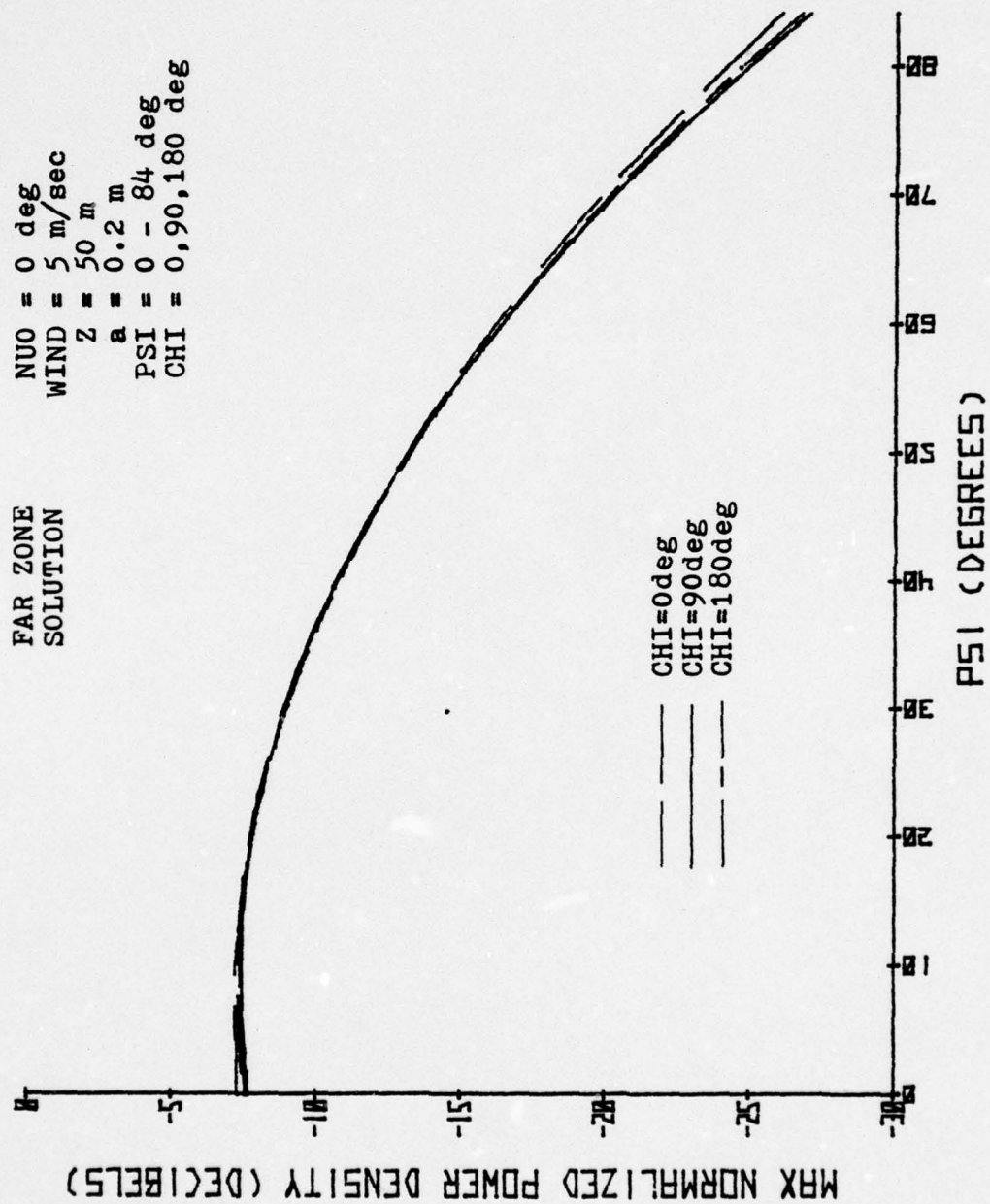


Figure 24

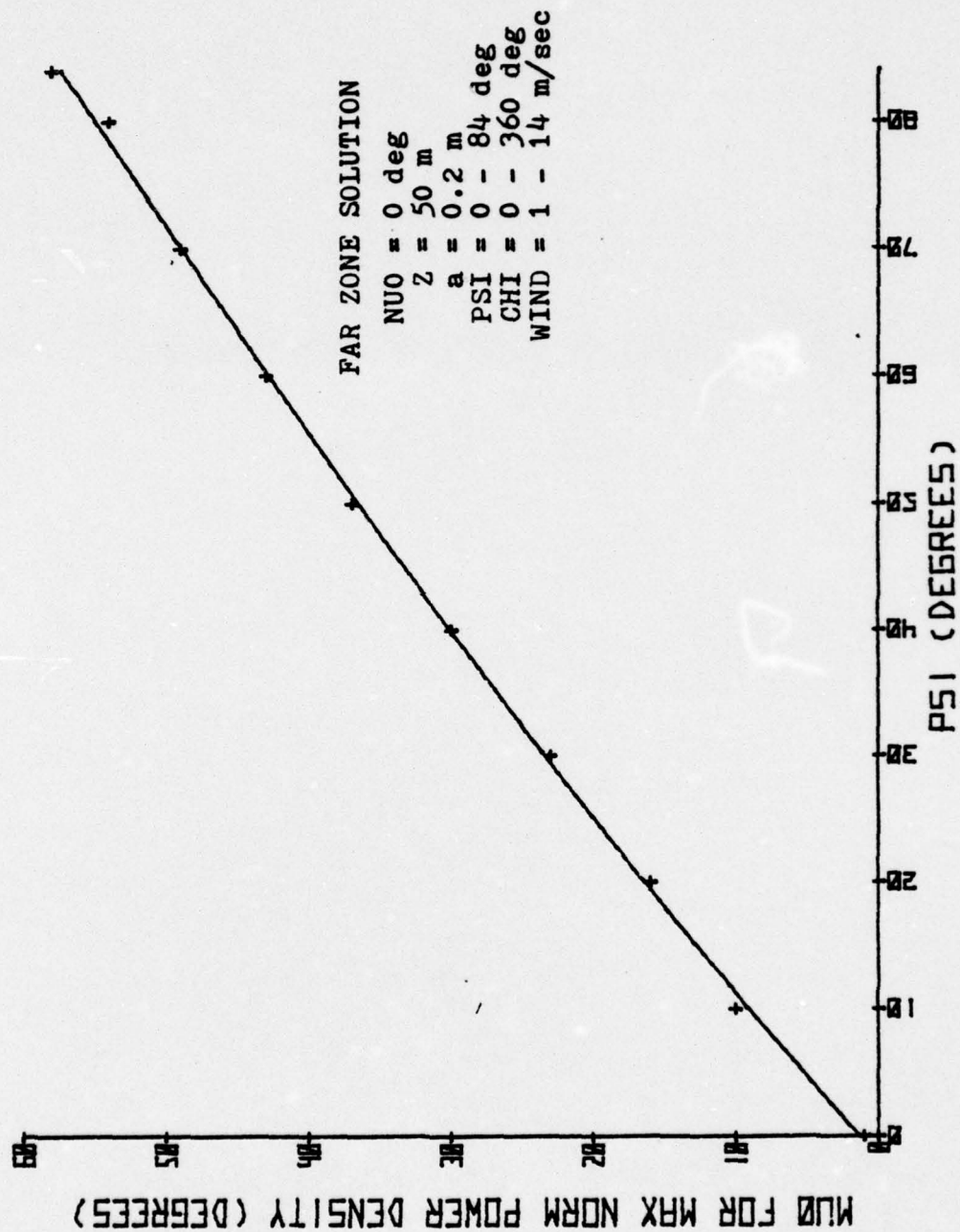


Figure 25

APPENDIX B

COMPUTER PROGRAM USAGE

To use the computer programs listed following this Appendix, first compute the ratio of the depth of the point of observation below the ocean surface to the beam radius at the surface facet, Z/a . Use the Near Zone program for $Z/a \leq 100$ and the Far Zone program for $Z/a > 100$. The Near Zone program may be used for any value of Z/a , but a slight computation time and computer storage increase will result.

The only other basic requirement is the selection of eight input variables. The variables are listed below along with the line from the respective program in which they appear.

| <u>VARIABLE</u> | <u>NEAR</u> <u>ZONE</u> | <u>FAR</u> <u>ZONE</u> |
|--|----------------------------|---------------------------|
| Depth, Z (m) | 470 | 440 |
| Beam radius, RAD (m) | 480 | 450 |
| Angle Nu_0 , NUORUN (deg) | 490 | 480 |
| Backward scattering coefficient, BS | 560 | 530 |
| Absorption coefficient, $AC \text{ (m}^{-1}\text{)}$ | 570 | 540 |
| Beam incidence angle, PSIRUN (deg) | 630 | 590 |
| Wind direction, CHI (rad) | 680 | 640 |
| Wind velocity, W (m/sec) | 710 | 670 |

Z, RAD, NUORUN, BS and AC are direct program inputs

made on the indicated lines. PSI, CHI and W are selected from DATA statements 430 and 440 in the Near Zone program and 420 and 430 in the Far Zone program. The variable locations in DATA statements were used for multiple variations of the parameters in a slightly different form of the program than presented which is easily set up with DO LOOPS. The entry variables may also be varied in this same manner.

Both programs must be combined with the common package of subroutines that are listed following the basic program listings.

The output from both listed programs appear on the next two pages.

NEAR ZONE COLLISION
 PSI=40.0 DEG
 MUO= 0.0 DEG
 DEP H= 10.00 M
 RAD= 0.2 M
 CFI= 0.0 DEG
 DIF= 2.37845
 VAL ID=1
 INVA ID=0
 WIND=14.3 M/SEC

| MUO | ALPHA | B | W: | NR | T | JACOB | PSLOPE | PRATIO | PWRDEN | MUO | VALIDITY | XL | ETA |
|-----|-------|-------|-------|-------|--------|--------|------------|------------|--------|-----|----------|-----|-----|
| 15 | 0.0 | 29.93 | 69.93 | 44.93 | 0.2999 | 0.5512 | 1.7628E-01 | 8.5333E-05 | -20.69 | 15 | 0 | 0.0 | 2.6 |
| 16 | 0.0 | 28.66 | 68.66 | 43.66 | 0.2866 | 0.4809 | 1.7185E-01 | 8.5333E-05 | -18.57 | 16 | 0 | 0.0 | 2.3 |
| 17 | 0.0 | 27.42 | 67.42 | 42.42 | 0.2742 | 0.4186 | 1.6750E-01 | 8.5333E-05 | -16.37 | 17 | 0 | 0.0 | 2.0 |
| 18 | 0.0 | 26.19 | 66.19 | 41.19 | 0.2619 | 0.3624 | 1.6407E-01 | 8.5333E-05 | -14.22 | 18 | 0 | 0.0 | 1.7 |
| 19 | 0.0 | 24.92 | 64.92 | 39.92 | 0.2492 | 0.3186 | 1.6077E-01 | 8.5333E-05 | -12.13 | 19 | 0 | 0.0 | 1.4 |
| 20 | 0.0 | 23.61 | 63.61 | 38.61 | 0.2361 | 0.2824 | 1.5752E-01 | 8.5333E-05 | -10.00 | 20 | 0 | 0.0 | 1.1 |
| 21 | 0.0 | 22.26 | 62.26 | 37.26 | 0.2226 | 0.2513 | 1.5440E-01 | 8.5333E-05 | -7.84 | 21 | 0 | 0.0 | 0.8 |
| 22 | 0.0 | 20.88 | 60.88 | 35.88 | 0.2088 | 0.2243 | 1.5142E-01 | 8.5333E-05 | -5.65 | 22 | 0 | 0.0 | 0.5 |
| 23 | 0.0 | 19.48 | 59.48 | 34.48 | 0.1948 | 0.2019 | 1.4859E-01 | 8.5333E-05 | -3.43 | 23 | 0 | 0.0 | 0.2 |
| 24 | 0.0 | 18.05 | 58.05 | 33.05 | 0.1805 | 0.1824 | 1.4597E-01 | 8.5333E-05 | -1.19 | 24 | 0 | 0.0 | 0.0 |
| 25 | 0.0 | 16.59 | 56.59 | 31.59 | 0.1659 | 0.1659 | 1.4357E-01 | 8.5333E-05 | 0.00 | 25 | 0 | 0.0 | 0.0 |
| 26 | 0.0 | 15.11 | 55.11 | 30.11 | 0.1511 | 0.1511 | 1.4137E-01 | 8.5333E-05 | 0.00 | 26 | 0 | 0.0 | 0.0 |
| 27 | 0.0 | 13.61 | 53.61 | 28.61 | 0.1361 | 0.1361 | 1.3935E-01 | 8.5333E-05 | 0.00 | 27 | 0 | 0.0 | 0.0 |
| 28 | 0.0 | 12.09 | 52.09 | 27.09 | 0.1209 | 0.1209 | 1.3750E-01 | 8.5333E-05 | 0.00 | 28 | 0 | 0.0 | 0.0 |
| 29 | 0.0 | 10.55 | 50.55 | 25.55 | 0.1055 | 0.1055 | 1.3588E-01 | 8.5333E-05 | 0.00 | 29 | 0 | 0.0 | 0.0 |
| 30 | 0.0 | 9.00 | 49.00 | 24.00 | 0.0900 | 0.0900 | 1.3448E-01 | 8.5333E-05 | 0.00 | 30 | 0 | 0.0 | 0.0 |
| 31 | 0.0 | 7.44 | 47.44 | 22.44 | 0.0744 | 0.0744 | 1.3329E-01 | 8.5333E-05 | 0.00 | 31 | 0 | 0.0 | 0.0 |
| 32 | 0.0 | 5.87 | 45.87 | 20.87 | 0.0587 | 0.0587 | 1.3231E-01 | 8.5333E-05 | 0.00 | 32 | 0 | 0.0 | 0.0 |
| 33 | 0.0 | 4.29 | 44.29 | 19.29 | 0.0429 | 0.0429 | 1.3155E-01 | 8.5333E-05 | 0.00 | 33 | 0 | 0.0 | 0.0 |
| 34 | 0.0 | 2.70 | 42.70 | 17.70 | 0.0270 | 0.0270 | 1.3098E-01 | 8.5333E-05 | 0.00 | 34 | 0 | 0.0 | 0.0 |
| 35 | 0.0 | 1.10 | 41.10 | 16.10 | 0.0110 | 0.0110 | 1.3059E-01 | 8.5333E-05 | 0.00 | 35 | 0 | 0.0 | 0.0 |
| 36 | 0.0 | -0.50 | 39.50 | 14.50 | 0.0050 | 0.0050 | 1.3037E-01 | 8.5333E-05 | 0.00 | 36 | 0 | 0.0 | 0.0 |
| 37 | 0.0 | -2.09 | 37.91 | 12.91 | 0.0009 | 0.0009 | 1.3031E-01 | 8.5333E-05 | 0.00 | 37 | 0 | 0.0 | 0.0 |
| 38 | 0.0 | -3.67 | 36.33 | 11.33 | 0.0003 | 0.0003 | 1.3040E-01 | 8.5333E-05 | 0.00 | 38 | 0 | 0.0 | 0.0 |
| 39 | 0.0 | -5.25 | 34.76 | 9.76 | 0.0000 | 0.0000 | 1.3064E-01 | 8.5333E-05 | 0.00 | 39 | 0 | 0.0 | 0.0 |

FAP TIME SLUT: 71
 PFI=60.0 DEG
 MUD=3.0 DEG
 DEPTH=50.00 M
 CH=0.0 DEG
 WTE=0.01121
 VALIC=0
 INVALID=0
 WIND=14.0 M/SEC

| MUD | ALPHA | R | WT | MR | T | JACOB | DAPEA | PSLOPE | PRATTN | PHRDEN | MUC | VALIDITY | XI | ETA |
|-----|-------|--------|-------|-------|--------|--------|-------------|------------|------------|--------|-----|----------|------|-----|
| 15 | 0.0 | 29.93 | 68.93 | 44.93 | 0.2549 | 0.5512 | 2.2846E-04 | 1.7628E-01 | 7.9106E-08 | -3.04 | 15 | 0 | 0.00 | 7.4 |
| 16 | 0.0 | 28.32 | 68.32 | 43.32 | 0.3143 | 0.4609 | 2.1162E-04 | 2.3785E-01 | 1.2603E-07 | -3.02 | 16 | 0 | 0.00 | 2.2 |
| 17 | 0.0 | 26.66 | 66.66 | 42.66 | 0.3143 | 0.3841 | 1.8128E-04 | 3.1207E-01 | 1.9531E-07 | -3.03 | 17 | 0 | 0.00 | 2.2 |
| 18 | 0.0 | 25.92 | 65.92 | 42.11 | 0.3698 | 0.2624 | 1.6944E-04 | 5.3525E-01 | 2.9479E-07 | -3.04 | 18 | 0 | 0.00 | 2.0 |
| 19 | 0.0 | 23.11 | 63.11 | 41.23 | 0.3975 | 0.2143 | 1.5934E-04 | 9.0404E-01 | 4.4079E-07 | -3.05 | 19 | 0 | 0.00 | 1.5 |
| 20 | 0.0 | 21.23 | 61.23 | 40.26 | 0.4251 | 0.1730 | 1.4894E-04 | 1.5404E-01 | 6.5350E-07 | -3.06 | 20 | 0 | 0.00 | 1.3 |
| 21 | 0.0 | 19.26 | 59.26 | 39.20 | 0.4521 | 0.1375 | 1.3892E-04 | 1.1812E-01 | 1.4323E-06 | -3.07 | 21 | 0 | 0.00 | 1.3 |
| 22 | 0.0 | 17.20 | 57.20 | 38.07 | 0.4788 | 0.1012 | 1.2902E-04 | 1.1544E-01 | 2.0907E-06 | -3.08 | 22 | 0 | 0.00 | 1.0 |
| 23 | 0.0 | 15.04 | 55.04 | 36.94 | 0.5048 | 0.0817 | 1.2002E-04 | 2.5767E-01 | 4.183E-06 | -3.09 | 23 | 0 | 0.00 | 1.0 |
| 24 | 0.0 | 12.78 | 52.78 | 35.71 | 0.5300 | 0.0592 | 1.1102E-04 | 3.2231E-01 | 6.1994E-06 | -3.10 | 24 | 0 | 0.00 | 0.7 |
| 25 | 0.0 | 10.41 | 50.41 | 34.48 | 0.5542 | 0.0433 | 1.0174E-04 | 3.8914E-01 | 8.2259E-06 | -3.11 | 25 | 0 | 0.00 | 0.4 |
| 26 | 0.0 | 7.92 | 47.92 | 33.25 | 0.5775 | 0.0242 | 9.0177E-05 | 4.4875E-01 | 1.2462E-05 | -3.12 | 26 | 0 | 0.00 | 0.2 |
| 27 | 0.0 | 5.28 | 45.28 | 32.03 | 0.5995 | 0.0105 | 8.0677E-05 | 4.9488E-01 | 1.4460E-05 | -3.13 | 27 | 0 | 0.00 | 0.0 |
| 28 | 0.0 | 2.58 | 42.58 | 30.71 | 0.6202 | 0.0019 | 7.0383E-05 | 4.6105E-01 | 1.4817E-05 | -3.14 | 28 | 0 | 0.00 | 0.0 |
| 29 | 0.0 | -3.30 | 39.70 | 29.56 | 0.6395 | 0.0022 | 6.0102E-05 | 3.8715E-01 | 1.3788E-05 | -3.15 | 29 | 0 | 0.00 | 0.0 |
| 30 | 0.0 | -6.42 | 36.56 | 28.28 | 0.6573 | 0.0022 | 5.0102E-05 | 2.8715E-01 | 1.324E-05 | -3.16 | 30 | 0 | 0.00 | 0.0 |
| 31 | 0.0 | -9.74 | 33.28 | 26.96 | 0.6880 | 0.0032 | 4.0102E-05 | 1.8715E-01 | 1.246E-05 | -3.17 | 31 | 0 | 0.00 | 0.0 |
| 32 | 0.0 | -13.07 | 30.28 | 25.63 | 0.7078 | 0.0042 | 3.0102E-05 | 9.3898E-02 | 1.1555E-05 | -3.18 | 32 | 0 | 0.00 | 0.0 |
| 33 | 0.0 | -16.49 | 27.31 | 24.31 | 0.7117 | 0.0042 | 2.0102E-05 | 8.3898E-02 | 1.0475E-05 | -3.19 | 33 | 0 | 0.00 | 0.0 |
| 34 | 0.0 | -19.91 | 24.31 | 22.98 | 0.7206 | 0.0042 | 1.0102E-05 | 7.3898E-02 | 9.3898E-06 | -3.20 | 34 | 0 | 0.00 | 0.0 |
| 35 | 0.0 | -23.33 | 21.33 | 21.65 | 0.7277 | 0.0042 | 0.0102E-05 | 6.3898E-02 | 8.3898E-06 | -3.21 | 35 | 0 | 0.00 | 0.0 |
| 36 | 0.0 | -26.75 | 18.33 | 20.33 | 0.7328 | 0.0042 | -0.0102E-05 | 5.3898E-02 | 7.3898E-06 | -3.22 | 36 | 0 | 0.00 | 0.0 |
| 37 | 0.0 | -30.17 | 15.33 | 19.03 | 0.7359 | 0.0042 | -0.0202E-05 | 4.3898E-02 | 6.3898E-06 | -3.23 | 37 | 0 | 0.00 | 0.0 |
| 38 | 0.0 | -33.59 | 12.33 | 17.75 | 0.7359 | 0.0042 | -0.0302E-05 | 3.3898E-02 | 5.3898E-06 | -3.24 | 38 | 0 | 0.00 | 0.0 |
| 39 | 0.0 | -37.01 | 9.33 | 16.48 | 0.7359 | 0.0042 | -0.0402E-05 | 2.3898E-02 | 4.3898E-06 | -3.25 | 39 | 0 | 0.00 | 0.0 |

NEAR ZONE COMPUTER PROGRAM

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460
470
480

NEAR ZONE SOLUTION OF THE TIME-AVERAGE POWER PROBABILITY
DISTRIBUTION BELOW THE OCEAN OF A LASER BEAM INCIDENT ON SURFACE

NCMENCLATURE:

PSI=ANGLE OF INCIDENT BEAM WITH RESPECT TO VERTICAL AXIS
CHI=ANGLE OF WIND WITH RESPECT TO THE INCIDENT BEAM IN
THE XY-PLANE
ALPHA=ANGLE OF MAX WAVE SLOPE WITH RESPECT TO THE INCIDENT
BEAM IN THE XY-PLANE
B=MAX SLOPE OF WAVE FACET
NU=NUO=ANGLE BETWEEN REFRACTED RAY PROJECTED ONTO XY-PLANE
AND THE Y AXIS
MU=MUO=ANGLE BETWEEN REFLECTED RAY AND THE Z AXIS
Z=DEPTH BELOW OCEAN SURFACE (METERS)
RAD=INCIDENT RAY SPOT RADIUS AT WAVE FACET (MINOR AXIS OF
ELLIPSE WHEN PSI IS NON-ZERO)
BB=MAJOR AXIS OF SPOT INCIDENT ON WAVE FACET
WI=WIND VELOCITY (M/SEC)
WI=RAY INCIDENCE ANGLE
WR=RAY REFRACTION ANGLE
LTF=DIFFUSE TRANSMITTANCE FUNCTION
BS=BACKWARD SCATTERING COEFFICIENT
AC=ABSORPTION COEFFICIENT
T=FRSNEL'S TRANSMITTANCE FUNCTION
AN=INDEX OF REFRACTION
PSLOPE=SLOPE PROBABILITY FUNCTION
DAREA=AREA INCREMENT
JACOB=JACOBIAN
PWRDEN=POWER DENSITY (DECIBELS)

INTEGER VALID
REAL NUO,JMUO,JACOB,MUO,NUORUN
REAL MUU,MUL,NUU,NUL
COMMON MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,BB,JK
DIMENSION W(4),CHI(3),NCHI(3),RG(4),JMUC(90),PWRDEN(90),PSIN
1(110)

INPUT PARAMETERS FOR THIS RUN

DATA W,CHI,NCHI/14.,10.,5.,1.,0.,1.5708,3.14159,0.90,180/
DATA PSIN/0.,10.,20.,30.,40.,50.,60.,70.,80.,84./
DATA ETAL,ETAZ,X11,X12/.577350269,-.577350269,-.577350269,-.5773502
169/
Z = 10.
RAD = .2


```

C 120 INT = 15
C 125 CC 195 J=INT,90
      JMU0(J) = FLOAT(J)
      MLC = JMU0(J)/57.2958
      JK = J
      JJ = J-6
C
C      ALPHA,B,WI,WR ARE COMPUTED IN SUBROUTINE ANGLES:
C      CALL ANGLES (ALPHA,B,WI,WR)
C      REFS 2,3&4
C      T=FRESNEL'S TRANSMITTANCE
C      REFS 9&10
C      TPER = (2.*COS(WI)*SIN(WR)/SIN(WI+WR))**2
C      TPAR = (2.*COS(WI)*SIN(WR)/SIN(WI+WR)*COS(WI-WR))**2
C      T = (TPAR+TPER)/2.
C
C      JACCBIAN COMPUTED IN SUBROUTINE JACO
C      CALL JACO (ALPHA,B,JACCB)
C      REF 3: EQNS 48-52
C
C      PSLCPE=SLOPE PROBABILITY FUNCTION
C      REF 5: EQNS 5-9,12-18
C      SIGC = SQRT(.003+1.92E-03*W(I))
C      SIGU = SQRT(3.16E-03*W(I))
C      C21 = .01-.0086*W(I)
C      C23 = .04-.033*W(I)
C      C40 = .4
C      C42 = .12
C      C44 = .23
C      AP = ALPHA-CHI(L)
C      ZXP = TAN(B)*SIN(AP)
C      ZYP = TAN(B)*COS(AP)
C      XI = ZXP/SIGC
C      ETA = ZYP/SIGU
C      FACT1 = (1./2.*PI*(SIGC*SIGU))*EXP(-.5*(XI**2+ETA**2))
C      FACT2 = 1.-.5*C21*(XI**2-1.)*ETA-(1./6.)*C03*(ETA**3-3.*ETA)
C      FACT3 = (1./24.)*C40*(XI**4-6.*XI**2+3.)*.25*C22*(XI**2-1.)*(ETA**
12-1.)
C      FACT4 = (1./24.)*C04*(ETA**4-6.*ETA**2+3.)
C      FLSLOPE = FACT1*(FACT2+FACT3+FACT4)
C
C      VALIDITY CHECK (VALID=1; INVALID=0) ON SLCPE PROBABILITY
C      AXI = ABS(XI)
C      AETA = ABS(ETA)
C      VALID = 1
C      IF ((AXI.GT.2.5).OR.(AETA.GT.2.5)) VALID=0

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```

CE = SQRT(BB**2*CTNU**2-(BB**2/RAD**2)*(Y1**2-BB**2+X1*CTNU*(X1*CT
1N1-2.*Y1)))
CC = CTNU**2+BB**2/RAD**2
X4 = (CA-(SIN(NUO))/ABS(SIN(NUO)))*CB)/CC
X5 = (CA+(SIN(NUO))/ABS(SIN(NUO)))*CB)/CC
Y4 = BB*SQRT(1.-X4**2/RAD**2)
Y4SG = (X4-X1)*CTNU+Y1
Y4 = SIGN(Y4,Y4SG)
Y5 = BB*SQRT(1.-X5**2/RAD**2)
Y5SG = (X5-X1)*CTNU+Y1
Y5 = SIGN(Y5,Y5SG)
GC TO 165

```

C NORMAL INCIDENCE, OUTSIDE BEAM:

```

155 X1 = 0.
Y1 = -2.*TAN(MUO)
X21 = -(RAD*SQRT(Y1**2-RAD**2)/Y1)
X22 = -X21
Y21 = -RAD*SQRT(1.-X21**2/RAD**2)
Y22 = -RAD*SQRT(1.-X22**2/RAD**2)
IF (NUO.NE.0.) GO TO 160
X4 = 0.
X5 = 0.
Y4 = 0.
Y5 = 0.
Y4SG = -RAD
Y5SG = RAD
GC TO 165

```

```

16C CTNU = COTAN(MUO)
CNA = -Y1*CTNU
CNB = SQRT(RAD**2*CTNU**2-Y1**2+RAD**2)
CNC = CTNU**2+1.
X4 = (CNA-(SIN(NUO))/ABS(SIN(NUO)))*CNB)/CNC
X5 = (CNA+(SIN(NUO))/ABS(SIN(NUO)))*CNB)/CNC
Y4 = BB*SQRT(1.-X4**2/RAD**2)
Y4SGN = X4*CTNU+Y1
Y4 = SIGN(Y4,Y4SGN)
Y5 = BB*SQRT(1.-X5**2/RAD**2)
Y5SGN = X5*CTNU+Y1
Y5 = SIGN(Y5,Y5SGN)
B1 = SQRT((Y4-Y1)**2+(X4-X1)**2)
B2 = SQRT((Y5-Y1)**2+(X5-X1)**2)
165 DNU1 = ARCCOS((Y1*(Y1-X21))/B1)
DNU2 = ARCCOS((Y1*(Y1-X22))/B2)
11#2+(X1-X21)**2)/SQRT((X1**2+Y1**2)*(Y1-Y21
11#2+(X1-X22)**2))/SQRT((X1**2+Y1**2)*(Y1-Y22
MLL = ATAN(B1/Z)
MLL = ATAN(B2/Z)
NLL = NUO+DNU1
NLL = NUO-DNU2

```



```

C      CCNVERT ALPHA,B,WI,WR FROM RADIANS TO DEGREES
RCEG = 57.2958
ALPHAD = ALPHA*RDEG
BD = B*RDEG
WIC = WI*RDEG
WRD = WR*RDEG
WRITE (6,190) J,ALPHAD,BD,WIC,WRC,T,JACOB,PSLCPE,PRATIO,PWRDEN(J),
1J,VALIC,AXI,AEIA
150 FCRMAT (1X,12,1X,F7.2,1X,F7.2,1X,F7.2,1X,F7.4,1X,F7.4,1X,1
1PE11.4,1X,1PE11.4,1X,OPF8.2,2X,12,5X,12,3X,F4.1,1X,F4.1)
C      LCCP EXIT FOR LOW SLOPE PROBABILITY
IF ((PSLOPE.LT.1.E-20).AND.(PWRDEN(J).LT.PWRDEN(JJ)).AND.(J.GT.5))
1 GC TO 200
155 CCNTINUE
C      200 CCNTINUE
END

```

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FAR ZONE COMPUTER PROGRAM

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FAR ZONE SOLUTION OF THE TIME-AVERAGE POWER PROBABILITY
DISTRIBUTION BELOW THE OCEAN OF A LASER BEAM INCIDENT
ON THE SURFACE

NCMENCLATURE:

PSI=ANGLE OF INCIDENT BEAM WITH RESPECT TO VERTICAL AXIS
CHI=ANGLE OF WIND WITH RESPECT TO THE INCIDENT BEAM IN
THE XY-PLANE
ALPHA=ANGLE OF MAX WAVE SLOPE WITH RESPECT TO THE INCIDENT
BEAM IN THE XY-PLANE
B=MAX SLOPE OF WAVE FACET
NU=NUO=ANGLE BETWEEN REFRACTED RAY PROJECTED ONTO XY-PLANE
AND THE Y AXIS
MU=MUO=ANGLE BETWEEN REFRACTED RAY AND THE Z AXIS
Z=DEPTH BELOW OCEAN SURFACE (METERS)
RAD=INCIDENT RAY SPOT RADIUS AT WAVE FACET (MINOR AXIS OF
ELLIPSE WHEN PSI IS NON-ZERO)
BB=MAJOR AXIS OF SPOT INCIDENT ON WAVE FACET
W=WIND VELOCITY (M/SEC)
WI=RAY INCIDENCE ANGLE
WR=RAY REFRACTION ANGLE
LTF=DIFFUSE TRANSMITTANCE FUNCTION
BS=BACKWARD SCATTERING COEFFICIENT
AC=ABSORPTION COEFFICIENT
T=FFRESNEL'S TRANSMITTANCE FUNCTION
AN=INDEX OF REFRACTION
PSLOPE=SLOPE PROBABILITY FUNCTION
DAREA=AREA INCREMENT
JACOB=JACOBIAN
PWRDEN=POWER DENSITY (DECIBELS)

INTEGER VALID
REAL NUO, JMUO, JACOB, MUO, NUORUN
COMMON MUO, NUO, PSI, AN, PSIRUN, NUORUN, PI, Z, RAD, BB, JK
DIMENSION W(3), CHI(3), NCHI(3), RG(4), JMUC(90), PWRDEN(90), PSIN
1(1C)

INFLUT PARAMETERS:
DATA W, CHI, NCHI/14.,10.,5.,0.,1.5708,3.14159,0.90,180/
DATA PSIN/0.,10.,20.,30.,40.,50.,60.,70.,80.,84./
Z = 50.
RAD = .2
PI = 3.141592654
AN = 1.33
NUORUN = 0.

CC

CC


```

970 JML0(J) = FLOAT(J)
980 MLC = JMU0(J)/57.2958
990 JK = J
1000 J = J-6
1010
1020 ALPHA,B,WI,WR ARE COMPUTED IN SUBROUTINE ANGLES:
1030 REFS 2,3&4
1040 CALL ANGLES (ALPHA,B,WI,WR)
1050
1060 T=FRESNEL'S TRANSMITTANCE
1070 REFS 9&10
1080 TPER = (2.*COS(WI)*SIN(WR))/SIN(WI+WR)**2
1090 TPAR = (2.*COS(WI)*SIN(WR))/SIN(WI+WR)*COS(WI-WR)**2
1100 T = (TPAR+TPER)/2.
1110
1120 JACOBIAN COMPUTED IN SUBROUTINE JACO
1130 REF 3: EQNS 48-52
1140 CALL JACO (ALPHA,B,JACOB)
1150
1160 AREA INCREMENT (DAREA) COMPUTED IN SUBROUTINE AREA
1170 REF 3: EQNS 13,14,21-24,38,39,79,80
1180 CALL AREA (DAREA)
1190
1200 FSLCPE=SLOPE PROBABILITY FUNCTION
1210 REF 5: EQNS 5-9,12-18
1220 SIGC = SQRT(.003+1.92E-03*W(I))
1230 SIGU = SQRT(3.16E-03*W(I))
1240 C21 = .01-.0086*W(I)
1250 C23 = .04-.033*W(I)
1260 C40 = .4
1270 C22 = .12
1280 C24 = .23
1290 AP = ALPHA-CHI(L)
1300 ZXP = TAN(B)*SIN(AP)
1310 ZY = TAN(B)*COS(AP)
1320 XI = ZXP/SIGC
1330 ETA = ZYP/SIGU
1340 FACT1 = (1./(2.*PI*SIGC*SIGU))*EXP(-.5*(XI**2+ETA**2))
1350 FACT2 = (1.-.5*C21*(XI**2-1.)*ETA-(1./6.)*C03*(ETA**3-3.*ETA)
1360 FACT3 = (1./24.)*C40*(XI**4-6.*XI**2+3.)*+.25*C22*(XI**2-1.)*(ETA**
1370 12-1.)
1380 FACT4 = (1./24.)*C04*(ETA**4-6.*ETA**2+3.)
1390 FSLCPE = FACT1*(FACT2+FACT3+FACT4)
1400
1410 VALIDITY CHECK (VALID=1; INVALID=0) ON SLCPE PROBABILITY
1420 AXI = ABS(XI)
1430 AETA = ABS(ETA)
1440 VALID = 1

```

```

1450 IF ((AXI.GT.2.5).OR.(AETA.GT.2.5)) VALID=0
1460 PRATIO=POWER DENSITY RATIO
1470 REFS 2,3&4
1480 PRATIO = ABS(DTF*TAN(B)*((1./CCS(B))**2)*CCS(WI)*(1./CCS(WR)))*PS
1490 1LLCPE*CAREA/JACOB)
1500
1510 PWRDEN(J) = 10.*ALOG10(PRATIO*Z**2)
1520 REF 2: PGS 65-67; FIG 34
1530
1540 CLTPUT INSTRUCTIONS
1550 CCNVERT ALPHA,B,WI,WR FROM RADIANS TO DEGREES
1560 RCEG = 57.2958
1570 ALPHAD = ALPHA*RDEG
1580 BD = B*RDEG
1590 WID = WI*RDEG
1600 WRTE = WR*RDEG
1610 WRTE (6,130) J,ALPHAD,BD,WID,WRD,T,JACOB,CAREA,PSLCPE,PRATIO,PWRD
1620 LEN(J); J,VALID,AXI,AETA
1630 13C FCRMAT 11X,12,1X,F7.2,1X,F7.2,1X,F7.2,1X,F7.4,1X,1
1640 1PE11.4,1X,1PE11.4,1X,1PF8.2,2X,12,5X,12,3X,F4.1,1X,F4.1
1650 2)
1660
1670 LCCP EXIT FOR LOW SLOPE PROBABILITY
1680 IF ((PSLCPE.LT.1.E-40).AND.(PWRDEN(J).LT.PWRDEN(JJ)).AND.(J.GT.5))
1690 1 GC TO 140
1700 135 CCNTINUE
1710
1720 14C CCNTINUE
1730 END
1740

```

COMMON COMPUTER SUBROUTINES

```

10 SUBROUTINE ANGLES (ALPHA,B,WI,WR)
20 SUBROUTINE ANGLES SOLVES FOR ALPHA,B,WI&WR GIVEN PSI,NUO,MUO&AN
30
40 EXTERNAL FINDAB,FINDB
50 REAL NUC,MUO,NUORUN
60 COMMON MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,EB,JK
70 DIMENSION Q(2),Y(1)
80 IF (PSIRUN.EQ.0.) GO TO 230
90 IF (NUORUN.EQ.0.) GO TO 215
100 SOLUTION FOR PSI&NUO NON-ZERO:
110 ALPHA & B FOR A GIVEN MUO & NUO ARE OBTAINED BY SOLVING
120 SIMULTANEOUS NON-LINEAR EQNS 82&83 OF REF 3 USING IBM SOURCE
130 LIB SUBROUTINE (NLNSYS)
140
150 FIRST ESTABLISH INITIAL GUESS FOR NLNSYS
160 IF ((JK.EQ.1).OR.((JK.EQ.15).AND.(PSIRUN.GT.30.))) GO TO 205
170 Q(1) = ALPHA
180 Q(2) = B
190 GC TO 210
200 G(1) = .1
210 G(2) = .1
220 CALL NLNSYS (2,10,4,ISING,1,FINDAB,Q)
230 ALPHA = Q(1)
240 B = Q(2)
250 WI = ARCOS(COS(PSI)*COS(B)-SIN(PSI)*COS(ALPHA)*SIN(B))
260 WR = ARCOS((1./AN)*SQRT(COS(WI)**2+AN**2-1.))
270 GC TO 235
280
290 SOLUTION FOR PSI NON-ZERO & NUO=0.0:
300 ALPHA = 0.
310 B SOLVED FROM NON-LINEAR EQN 82 OF REF 3 USING IBM SOURCE LIB
320 SUBROUTINE (NLNSYS): FIRST ESTABLISH INITIAL B GUESS
330 IF ((JK.EQ.1).OR.((JK.EQ.15).AND.(PSIRUN.GT.30.))) GO TO 220
340 Y(1) = B
350 GC TO 225
360 Y(1) = .1
370 CALL NLNSYS (1,10,4,ISING,1,FINDB,Y)
380 B = Y(1)
390 WI = ARCOS(COS(PSI)*COS(B)-SIN(PSI)*SIN(B))
400 WR = ARCOS((1./AN)*SQRT(COS(WI)**2+AN**2-1.))
410 GC TO 235
420
430 SOLUTION FOR PSI=0.0:
440 ALPHA = NUO
450 B = ATAN((AN*SIN(MUO))/(AN*COS(MUO)-1.))
460 WI = B
470
480

```


C 235 $WR = \text{ARSIN}((1./AN)*\text{SIN}(B))$
 RETURN
 END

490
500
510
520

```

C C C C C
SLROUTINE FINDAB (Q,F,K)
SLROUTINE USED BY NLNSYS FCR COMPUTING THE (ALPHA&B) PAIR
      Q(1)=ALPHA
      Q(2)=B
REAL MLO,NUO,NUORUN
COMMON MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,BB,JK
DIMENSION Q(2)
IF (K.EQ.1) GO TO 305
GC TO 310
C 305 F = -COS(MUO)+(1./AN)*(COS(PSI)+(SIN(PSI)*CCS(Q(1))*SIN(Q(2)))-COS(
      1)PSI)*COS(Q(2))+SQRT((SIN(PSI)*COS(Q(1))*SIN(Q(2)))-COS(PSI)*COS(Q(2
      2)))*2+AN**2-1.))*COS(Q(2)))
      GC TO 315
C 310 F = -COTAN(NUO)+COTAN(Q(1))-SIN(PSI)/(SIN(Q(1))*SIN(Q(2)))*(SIN(PSI
      1)*COS(Q(1))*SIN(Q(2))-COS(PSI)*CCS(Q(2))+SQRT((SIN(PSI)*COS(Q(1))*
      2SIN(Q(2))-COS(PSI)*COS(Q(2)))*2+AN**2-1.))
C 315 RETURN
      END

```

```

C
C
C
SUBROUTINE FINDB (Y,F,K)
SUBROUTINE FINDB USED BY NLNSYS FCR COMPUTING B; B=Y(1)
REAL MUO,NUO,NUORUN
COMMON MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,BB,JK
DIMENSION Y(1)
SPSI = SIN(PSI)
CPSI = COS(PSI)
C
C
F = -COS(MUO)+(1./AN)*(CPSI*SIN(Y(1)))-CPSI*COS(Y(1))+SQRT((S
1PSI*SIN(Y(1)))-CPSI*COS(Y(1)))**2+AN**2-1))*COS(Y(1))
RETURN
END
C

```

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110
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```

10 SLROUTINE JACO (ALPHA,B,JACOB)
20
30 SLROUTINE JACO COMPUTES THE JACOBIAN FOR NORMAL & OBLIQUE
40 INCIDENCE.
50
60 REAL JACOB,MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,BB,JK
70 CCMCN MUO,NUO,PSI,AN,PSIRUN,NUORUN,PI,Z,RAC,BB,JK
80 IF (PSIRUN.EQ.0.) GO TO 505
90
100 JACOBIAN FOR OBLIQUE INCIDENCE:
110 AKK = SQRT((SIN(PSI)*COS(ALPHA)*SIN(B)-COS(PSI)*COS(B))*2+AN**2-1
120 1.)
130 SK = (SIN(PSI)*COS(ALPHA)*SIN(B)-COS(PSI)*COS(B))/AKK
140 CKCA = ((SIN(PSI)*SIN(ALPHA)*SIN(B)-COS(PSI)*COS(B)-SIN(PSI)*COS(
150 1ALPHA)*SIN(B))/AKK)-SIN(PSI)*SIN(ALPHA)*SIN(B)
160 DKCB = ((SIN(PSI)*COS(ALPHA)*SIN(B)-COS(PSI)*COS(B))*2+(SIN(PSI)*CO
170 1S(ALPHA)*COS(B)+COS(PSI)*SIN(B))/AKK)+(SIN(PSI)*COS(ALPHA)*COS(B)
180 2+CCS(PSI)*SIN(B))
190 D = ((SK*SIN(ALPHA)*SIN(B))*2+(SK*COS(ALPHA)*SIN(B)-SIN(PSI))*2)
200 1*SQRT(AN**2-(COS(PSI)+SK*COS(B))*2)
210 JACOB = ABS((1./D)*((COS(B)*SIN(ALPHA)*SIN(PSI)*DKCA)*(DKCB*SIN(B)
220 1+SK*COS(B))+DKCB*COS(B)-SK*SIN(B))*((SK*SIN(B))*2-(SIN(PSI)*SIN(
230 2B))*DKDA*SIN(ALPHA)+SK*COS(ALPHA))))
240 GC TO 510
250
260 JACOBIAN FOR NORMAL INCIDENCE:
270 505 AK = -CCS(B)+SQRT(COS(B))*2+AN**2-1.)
280 DKCB = SIN(B)*(1.-(COS(B)/SQRT(COS(B))*2+AN**2-1.))
290 JACOB = ABS((1./SQRT(AN**2-(1.+AK*COS(B))*2))*DKCB*COS(B)-AK*SIN
300 1(B)))
310
320 C 510 RETURN
330 ENC

```



```

1-X21)**2))
C4=(Y1*(Y1-X22)+X1*(X1-X22))/SQRT((X1**2+Y1**2)*((Y1-Y21)**2+(X1
1-X21)**2))
IF (C3.GT.1.0) C3 = 1.0
IF (C4.GT.1.0) C4 = 1.0
LAREA = ((PI/4.)*(ATAN(C1)-ATAN(C2)))*(ARCOS(C3)+ARCOS(C4))
GC TO 640
C
C AREA FOR NORMAL INCIDENCE:
625 X1 = 0.
Y1 = -Z*TAN(MUO)
X21 = -(RAD*Y1*SQRT(Y1**2-RAD**2))/Y1**2
X22 = -X21
Y21 = -RAD*SQRT(1.-X21**2/RAD**2)
Y22 = -RAD*SQRT(1.-X22**2/RAD**2)
IF (MUO.NE.0.) GO TO 630
X4 = 0.
X5 = 0.
Y4 = -RAD
Y5 = RAD
GC TO 635
630 CTNU = COTAN(MUO)
CNA = -Y1*CTNU
CNB = SQRT(ABS(RAD**2*CTNU**2-Y1**2+RAD**2))
CNC = CTNU**2+1.
X4 = (CNA-(SIN(MUO)/ABS(SIN(MUO))))*CNB)/CNC
X5 = (CNA+(SIN(MUO)/ABS(SIN(MUO))))*CNB)/CNC
Y4 = BB*SQRT(1.-X4**2/RAD**2)
Y4SGN = X4*CTNU+Y1
Y5 = SIGN(Y4,Y4SGN)
Y5SGN = BB*SQRT(1.-X5**2/RAD**2)
Y5SGN = X5*CTNU+Y1
Y5 = SIGN(Y5,Y5SGN)
C1 = SQRT((Y5-Y1)**2+(X5-X1)**2)/Z
C2 = SQRT((Y4-Y1)**2+(X4-X1)**2)/Z
C3 = (Y1*(Y1-Y21)+X1*(X1-X21))/SQRT((X1**2+Y1**2)*((Y1-Y21)**2+(X1
1-X21)**2))
C4 = (Y1*(Y1-Y22)+X1*(X1-X22))/SQRT((X1**2+Y1**2)*((Y1-Y21)**2+(X1
1-X21)**2))
IF (C3.GT.1.0) C3 = 1.0
IF (C4.GT.1.0) C4 = 1.0
LAREA = ((PI/4.)*(ATAN(C1)-ATAN(C2)))*(ARCCS(C3)+ARCCS(C4))
C
640 RETURN
END

```



```

C
      IF PARTIAL IS TOO SMALL, INCREASE TALLY.
      IF (ABS(PART(IITEMP)).EQ.0) GO TO 9
      IF (ABS(F/PART(IITEMP)).GT.1.0E20) TALLY=TALLY+1
      GO TO 10
9     TALLY=TALLY+1
10    CCNTINUE
      IF (TALLY.LE.(N-K)) GO TO 15
      FACTOR=FACTOR*10.0
      IF SURFACE IS TOO FLAT, THE SINGULARITY INDICATOR IS SET TC
      ZERO AND RETURN IS EXECUTED.
      IF (FACTOR.GT..5) GO TO 65
      GO TO 7
15   IF (K.LT.N) GO TO 20
      IF LAST PARTIAL IS ZERO, A SINGULARITY IS INDICATED AND A
      RETURN EXECUTED.
      IF (ABS(PART(IITEMP)).EQ.0.) GO TO 65
      CCE(K,N+1)=0
      KMAX=ITEMP
      GO TO 40
20   KMAX=PONTER(K,K)
      DERMATX=ABS(PART(KMAX))
      KPLUS=K+1
      GET INDEX FOR LARGEST PARTIAL IN K*TH EQUATION.
      DO 30 I=KPLUS,N
      JSUB=PONTER(K,I)
      TEST=ABS(PART(JSUB))
      IF (TEST.LT.DERMATX) GO TO 25
      DERMATX=TEST
      DEFINE PIVOT TO SWIVEL ABOUT THE VARIABLE WITH MAXIMUM PARTIAL
      WHEN WE GET TO THE NEXT EQUATION.
      PCNTER(KPLUS,I)=KMAX
      IF THIS PARTIAL IS BIGGER, WE HAVE A NEW MAXIMUM.
      KMAX=JSUB
      GO TO 30
25   PCNTER(KPLUS,I)=JSUB
30   CCNTINUE
      IF THAT PARTIAL IS 0, INDICATE A SINGULARITY AND RETURN.
      IF (ABS(PART(KMAX)).EQ.0) GO TO 65
      ISLB(K)=KMAX
      CCE(K,N+1)=0
      JSUB=PONTER(KPLUS,N)
      SAVE THESE CONSTANTS FOR FUTURE USE.
      CCE(K,JSUB)=-PART(JSUB)/PART(KMAX) NEW X(KMAX) VALUE
      GET PART OF EXPRESSION FOR THE NEW X(JSUB)
      CCE(K,N+1)=COE(K,N+1)+PART(JSUB)*X(JSUB)
      CCE(K,N+1)=(COE(K,N+1)-F)/PART(KMAX)+X(KMAX)
      IF N IS 1, WE HAVE OUR SOLUTION IN THE NEXT STEP WITHOUT ANY

```

NLNS1880
 NLNS1890
 NLNS1891
 NLNS1892
 NLNS1893
 NLNS1910
 NLNS1920
 NLNS1930
 NLNS1940
 NLNS1950
 NLNS1960
 NLNS1970
 NLNS1980
 NLNS1990
 NLNS2000
 NLNS2010
 NLNS2020
 NLNS2030
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 NLNS2100
 NLNS2110
 NLNS2120
 NLNS2130
 NLNS2140
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 NLNS2170
 NLNS2180
 NLNS2190
 NLNS2200
 NLNS2210
 NLNS2220
 NLNS2230
 NLNS2240
 NLNS2250
 NLNS2260
 NLNS2270
 NLNS2280
 NLNS2290
 NLNS2300
 NLNS2310
 NLNS2320
 NLNS2330

```

C      BACK-SUBSTITUTING.
C      X(KMAX)=COE(N,N+1)
C      FOR N GREATER THAN 1, WE PERFORM A FINAL BACK-SUBSTITUTION TO
C      GET OUR NEW X-VECTOR.
C      IF (N.GT.1) CALL BAKSUB(N,N,X)
C      IF (M.EQ.1) GO TO 50
C      DC 43 I=1,N
C      IF TEST FOR CONVERGENCE.
C      IF (ABS((TEMP(I)-X(I))/X(I)).GT.RELCON) GO TO 45
C      CCNTINUE
C      CCNVRG=CCNVRG+1
C      IF IT CONVERGES, RETURN WITH LAST VECTOR.
C      IF (CCNVRG.GE.3) GO TO 60
C      GC TO 50
C      45 CCNVRG=1
C      SAVE CURRENT X-VECTOR FOR TESTING WITH NEXT X-VECTOR.
C      DC 55 I=1,N
C      55 TEMP(I)=X(I)
C      IF M IS THE ITERATION LIMIT, RETURN.
C      GC TO 70
C      6C MAXIT=M
C      GC TO 70
C      65 ISING=1
C      70 RETURN
C      70 END

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NS2340
NLNS2350
NLNS2360
NLNS2370
NLNS2380
NLNS2390
NLNS2400
NLNS2410
NLNS2420
NLNS2430
NLNS2440
NLNS2450
NLNS2460
NLNS2470
NLNS2480
NLNS2490
NLNS2500
NLNS2510
NLNS2520
NLNS2530
NLNS2540
NLNS2550
NLNS2560
NLNS2570
NLNS2580

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